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3 **Systematic paleontology of Bartonian larger benthic foraminifera from Shahrekord region**
4
5 **in High Zagros, Iran**
6

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24

25 **Abstract**
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27
28 The Jahrum Formation is characterized by abundant benthic foraminifera in carbonate beds,
29 partly marly and dolomitic limestones at the Kuh-e-Soukhteh (Shahrekord region). This
30 formation covers a huge stretch of the Zagros Zone which is a part of the central Tethyan realm
31 during the Paleogene time. Biostratigraphic analysis of the larger benthic foraminifera is deduced
32 to distinguish one assemblage zone assigned to the late Middle Eocene (Bartonian). This new
33 biostratigraphic range is represented by the index fossil of *Rhabdorites malatyaensis* (Sirel) and
34 correlated with calcareous rocks in the Shiraz area (south Iran), Dhofar section (Oman), and
35 Socotra Island (Yemen). The Jahrum Formation is dominated by rich miliolids-agglutinated
36 foraminifera with rare small rotaliids and without *Nummulites* Lamarck and *Alveolina* d'Orbigny
37 indicating that the formation was deposited in a shallow water environment (nearshore lagoonal
38 zone) with low energy.
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3 **Keywords:** Bartonian, Iran, Jahrum Formation, Larger benthic foraminifera, Shahrekord,
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5 Systematics
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11 **Introduction**

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17 The separate continental blocks of the Iranian platform are jointed by ophiolitic units. The
18 geological zonation of this platform represents various sedimentary basins such as Alborz, Kopet
19 Dagh, Central Iran, Zagros, Sanandaj-Sirjan, Urumieh-Dokhtar, Lut Block, eastern Zone, and
20 Makran (Figure 1A). Because of different oil formations, various facies analysis and
21 stratigraphical studies have been carried out on the Zagros sedimentary basin, especially Jahrum
22 Formation (James and Wynd, 1965; Rahaghi, 1978, 1980, 1983; Kalantari, 1980; Stocklin and
23 Setudehnia, 1991; Vaziri-Moghaddam *et al.*, 2010). The Jahrum Formation consists of a
24 succession of thick layers to massive calcareous sedimentary rocks, thin and medium marls, and
25 dolomites with intercalation of yellow medium bedded limestones. The lower contact of this
26 formation with the Pabdeh Formation is faulted and the upper contact with the Asmari Formation
27 is unconformable. This work provides new biostratigraphic data on a Bartonian stratigraphic
28 interval of the Jahrum Formation in the study area. The foraminiferal assemblage allows
29 comparing with another biostratigraphic scheme for the adjacent areas (Fars area). The aim of
30 this paper is to examine in detail the microbiostratigraphy of the Jahrum Formation and to
31 introduce the systematic descriptions of larger foraminifera.
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54 **Geological setting**

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6 The geology of Shahrekord is controlled by numerous faults and consisted of three zones:
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8 Northeast (Zayandehrud), Southwest (Karun), and Central Zones (Zahedi and Rahmati Ilakhchi,
9
10 2006). The Central Zone (Z2) is a part of the high Zagros and located between the thrust faults
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12 (Saman - Fereidoon Shahr) (F1) and (Bazoft Fault) (F3) and divided into two smaller subzones
13
14 (Z2a) and (Z2b) by the main Zagros thrust fault (Figure 1B). These two sub-zones are located in
15
16 the Shahrekord region and consist of Cretaceous to Paleogene red clastic rocks, gray to cream
17
18 limestone, and marl correlating to the Jahrum, Pabdeh, and Asmari formations. The Z2b sub-
19
20 zone encompasses the high Zagros and is located between the main Zagros thrust fault (F2) and
21
22 the Bazoft thrust fault (F3) (Zahedi and Rahmati Ilakhchi, 2006). The stratigraphic section is
23
24 exposed on the roadside and located in Kuh-e-Soukhteh with geographical coordination of 32°
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26 $00' 00''$ N to $32^{\circ} 01' 21''$ N and $50^{\circ} 55' 51''$ E to $50^{\circ} 56' 80''$ E (Figure 1C).
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34 **Material and methods**

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39 The study is carried out based on thin sections of cemented carbonate rock. The stratigraphic
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41 thickness with 157 m, were measured from the sedimentary rocks in the Kuh-e-Soukhteh. A total
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43 of 80 rock samples were collected for the systematic study and precise identification of
44
45 foraminiferal species. The microscopic analysis of samples was carried out in the Geological Lab
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47 of Payame Noor University. The stratigraphic column and foraminiferal plate are done with
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49 Adobe Illustrator software. All samples are housed in Payame Noor University Laboratory.
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56 **Biostratigraphy**

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6 Based on the larger benthic foraminifera, the Eocene foraminiferal assemblage zones of Jahrum
7 Formation are established by James and Wynd (1965), Adams and Bourgeois (1967), and
8 Hottinger (2007) in the Zagros zone. In the current study, 18 species of benthic foraminifera
9 have been identified for the first time. The major taxa are found in the Jahrum Formation while
10 some taxa such as *Coskinolina* Stache, *Haymanella* Sirel, and *Austrotrillina* Parr have been
11 reported in other parts of Iran, such as the Alborz Mountains and eastern Iran (Babazadeh, 2003).
12 The benthic foraminiferal association is the same as the foraminiferal fauna of central Neo-
13 Tethys such as Shiraz area (south Iran) (Hottinger, 2007), Dhofar (Oman), and Socotra Island
14 (Yemen) (Serra-Kiel *et al.*, 2016). The Jahrum Formation is attributed to Bartonian based on the
15 concurrent range zone of the benthic foraminiferal association while in some previous works,
16 this formation was considered Ypresian and Lutetian (Early-Middle Eocene) in the study area
17 (Zahedi and Rahmati Ilakhchi, 2006; Babazadeh *et al.*, 2015). This association is collected from
18 the gray to cream-colored thin to thick-bedded limestones and argillaceous limestones and
19 occurred in one columnar section (Kuh-e-Soukhteh section). The distribution of selected benthic
20 foraminifera is shown in the stratigraphic section in Figure 2.
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45 **Systematic descriptions**

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50 Class Tubothalamea Pawlowski *et al.*, 2013

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53 Order Miliolida Delage and Herouard, 1896
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3 Family Hauerinidae Schwager, 1876
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6 Genus *Nurdanella* Özgen, 2000
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9 *Nurdanella boluensis* Özgen, 2000
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11 Figure 3A
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18 *Type species.* — *Nurdanella boluensis* Özgen, 2000
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20
21 *Description.* — In the record species, the axial and equatorial diameters of the test are 1.2 – 1.7
22 mm and 1.0 – 1.3 mm, respectively. The early chambers are arranged in a quinquelocoline
23 pattern and the later adult chambers are coiled in a planispiral mode.
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28 *Remarks.* — *Nurdanella boluensis* Özgen differs from *Nurdanella paleocenica* Sirel in its larger
29 size with high chambers. There are two chambers in the last whorl of the holotype of *Nurdanella*
30 Özgen (2000). Whereas, the *N. boluensis* has 5 to 6 chambers in the last whorl of the
31 microspheric form. The record species is close to *N. boluensis* due to its size. The measurement
32 of embryonic apparatus is not possible consequently the term of *confer* is used for this specimen.
33
34 The *N. boluensis* occurs in the Thanetian and Lutetian limestones of Turkey (Özgen, 2000; Sirel,
35 2013) but the biostratigraphic range of this species reaches to Bartonian stages in the study area.
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48 Family Hauerinidae Schwager, 1876 or Family Austrotrilinidae Loeblich and Tappan, 1987
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51 Genus *Austrotrillina* Parr, 1942
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54 *Austrotrillina eocaenica* Hottinger, 2007
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Figure 3B

Type species. — *Trillina howchini* Schlumberger, 1893

Description. — The *Austrorillina eocaenica* Hottinger is a true miliolid with an alveolar exoskeleton and thickened basal layer recorded from the Early ?- Lutetian–Priabonian of Shiraz in Iran (Hottinger, 2007), of Lampione Island near Sicily (southern Italy) (Di Carlo and Pignatti, 2009) of western Dhofar in Oman and Socotra Island in Yemen (Serra-Kiel *et al.*, 2016). The megalospheric forms of the *Austrorillina eocaenica* Hottinger show a subtriangular outline with a rounded margin in the axial section. The outline in the longitudinal section is ovate. The basal layer is thick. The maximum length in the axial sections for megalospheric forms is between 1.0 to 1.3 mm in the record species.

Remarks. — The morphology of Eocene austrotilinas is different from the Oligocene taxon. The *Austrorillina eocaenica* Hottinger differs from *Austrorillina paucialveolata* Grimsdale and *Austrorillina brunni* Marie (Oligocene forms) in a larger size, increased diameter of the proloculus, axial length in megalospheric forms, and the thicker basal layer. The *A. eocaenica* has the same kind of alveoli as *A. paucialveolata*. This species differs from *Austrorillina asmariensis* Adams (Oligocene form) in deduced diameter of axial length in megalospheric forms. In Oman and Iran, the *A. eocaenica* occurs in association with most of the conical agglutinated species as *Coskinolina perpera* (Hottinger, 2007). According to Adams (1968), the *Austrorillina* Parr ranges from Lower Oligocene to Lower Miocene (Loeblich and Tappan, 1987). Hottinger (2007) extended the biostratigraphic range of the *A. eocaenica* from Lutetian to Priabonian (SBZ 17-18). The recorded specimen of the *A. eocaenica* in the studied area is

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3 associated with *Neorhipidionina spiralis* Hottinger, *Nurdanella boluensis* Özgen,
4 *Quinqueloculina* sp. and other small miliolids. Its biostratigraphic range can be assigned to
5
6 Bartonian based on faunal association in the studied area.
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14 Family Soritidae Ehrenberg, 1839
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16 Genus *Macetadiscus* Hottinger, Serra- Kiel and Gallardo-Garcia, in: Serra- Kiel *et al.*, 2016,
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18 page 35 of 95
19

20 *Macetadiscus cf. incolumnatus* Hottinger, Serra-Kiel and Gallardo-Garcia, in: Serra- Kiel *et*
21
22 *al.*, 2016, page 36 of 95, figure 27 (1-27)
23

24
25 Figures 3C1 and 4G2
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29 *Type Species.* — *Macetadiscus incolumnatus* Hottinger, Serra- Kiel and Gallardo-Garcia
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32 *Description.* — In megalospheric forms of record species, the equatorial diameter is 2.6 mm
33
34 and the equatorial thickness is 0.3 mm. The microspheric forms are not observed in this study.
35
36 There are not endoskeletal structures such as pillars or partitions in the *Macetadiscus* Hottinger.
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39 *Remarks.* — The *Macetadiscus incolumnatus* Hottinger, Serra- Kiel and Gallardo-Garcia is
40
41 presented and figured as a new genus and new species in Serra- Kiel *et al.* (2016) for the first
42
43 time. The Late Paleocene Turkish specimen such as *Mardinella* Meric and Çoruh (type species:
44
45 *Orbitolites shirazensis* Rahaghi) was presented by Meric and Çoruh (1991) is very similar to
46
47 *Macetadiscus* Hottinger, Serra- Kiel and Gallardo-Garcia based on its morphostructure. On the
48
49 other hand, the figured specimen (pl. 18, figs. 10-12) of *Orbitolites* sp. in the Late Paleocene of
50
51 Shiraz area (Rahaghi, 1983) looks like *Macetadiscus* Hottinger, Serra- Kiel and Gallardo-Garcia
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55 Certainly. The *Mardinella* Meric and Çoruh could not be placed *O. shirazensis*. However, it
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3 seems more likely that the two genera (*Mardinella* and *Macetadiscus*) are synonymous. The
4 biostratigraphic range of the *Macetadiscus* cf. *incolumnatus* Hottinger, Serra- Kiel and Gallardo-
5 Garcia extends from the late Lutetian to the Priabonian (SBZ 16–20) according to Serra-Kiel *et*
6 *al.* (2016) and Nafarieh *et al.* (2019). In this work, the biostratigraphic range of recorded species
7 can be attributed to the Bartonian stage.
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18 Family Soritidae Ehrenberg, 1839

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20 Genus *Rhabdorites* Fleury, 1996

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23 *Rhabdorites malatyaensis* (Sirel, 1976)

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27 Figures 3C2, 3D and 3E2

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32 *Type species.* — *Rhapydionina malatyaensis* (Sirel), 1976

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35 *Description.* — There are 12-13 chambers in a uniserial arrangement with a diameter of 1.25
36 mm in the record species. The longitudinal and transversal diameters of the test are reached to
37 3.6 and 0.9 mm, respectively.
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43 *Remarks.* — The genus *Rhabdorites* Fleury differs from *Praerhapydionina* Van Wesseem in the
44 presence of multiple apertures. It differs from *Neotaberina* Hottinger in the absence of pillars in
45 the endoskeleton and more cylindrical morphology. The species of *Rhabdorites malatyaensis*
46 (Sirel) was described from Lutetian of Oman and Yemen under the taxon of *Rhabdorites*
47 *minimus* Henson (Serra-Kiel *et al.*, 2016). The biostratigraphic range of the *R. malatyaensis*
48 extends from Bartonian to Priabonian (Papazzoni and Sirotti, 1995; Serra-Kiel *et al.*, 2016;
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3 Nafarieh *et al.*, 2019). This species is associated with the last large *Orbitolites* spp. and *Malatyna*
4 *vicensis* Sirel (Sirel and Acar, 2008). According to Sirel (2003), *R. malatyaensis* is associated
5 with *Nummulites fabianii* (Perver) and *Nummulites biedai* Schaub in Arabil and Devely sections,
6 respectively. Therefore, the range of *R. malatyaensis* is shown to overlap that of *N. fabianii*. The
7 co-occurrence of *R. malatyaensis* with *Orbitolites minimus* Henson seems to draw the position of
8 the boundary between Middle Eocene and Upper Eocene (Hottinger, 2007). According to
9 Romero *et al.* (1999), the *R. malatyaensis* occurs in association with *Malatyna vicensis* Sirel and
10 Acar, in Igualada basin of northern Spain during the Bartonian. After Hottinger (2007), the *R.*
11 *malatyaensis* occurs in association with *Globoreticulina iranica* Rahaghi in the Shiraz area (west
12 Iran) and indicates the Bartonian age. In this work, the *Rhabdorites malatyaensis* (Sirel) is
13 observed in the Bartonian age. The biostratigraphical range of *R. malatyaensis* is particularly
14 important because most researchers proposed the Bartonian age based on the foraminiferal
15 associations.

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37 Family Soritidae Ehrenberg, 1839

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40 Genus *Praerhapydionina* Van Wessem, 1943

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43 *Praerhapydionina delicata* Henson, 1950

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Figure 3E1

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52 *Remarks.* — The *Praerhapydionina delicata* Henson differs from *Praerhapydionina huberi*
53 Henson by its small size (less than 2 mm). According to Sirel (2003), *P. delicata* and *P. huberi*

1
2
3 have the same biostratigraphic range. For the first time, the *P. delicata* was represented in
4 Oligocene from Iran and Iraq (Henson, 1950). After Fleury (1997), the biostratigraphic range of
5 *P. delicata* extends from the Middle to Late Eocene. This species is reported from Paleocene in
6 Iran (Rahaghi, 1983). It is also assigned to Bartonian-Rupelian by Hottinger (2007) and Serra-
7 Kiel *et al.* (2016). Based on foraminiferal association, the recorded species is occurred in the
8 Bartonian stage.
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21 Family Soritidae Ehrenberg, 1839

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23 Genus *Haymanella* Sirel, 1998

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27 ***Haymanella huberi*** (Henson, 1950)

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Figure 3F

Type species. — *Haymanella paleocaenica* Sirel, 1998

Description. — The *Haymanella* Sirel differs from the other praerhapydioninid genera by the presence of agglutinated grains in its porcelaneous wall (Sirel, 1998; Hottinger, 2007; Nafarieh *et al.*, 2019).

Remarks. — The *Haymanella huberi* (Henson) differs from the *Haymanella paleocaenica* Sirel in its shorter nepionic spiral stage in the megalospheric generation and its more complex foramina in the adult. The outline of the foramen is stellar with at least six branches of petaloid extensions. The *H. huberi* is about twice the size of *H. paleocaenica*. The *H. huberi* is known from Eocene deposits of Jahrum Formation (Iran) and Eocene deposits of Iraq, Turkey, and

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3 Oman (Henson, 1950; Sirel, 1998; Hottinger, 2007). According to Hottinger (2007) and Serra-
4
5 Kiel *et al.* (2016), the biostratigraphic range of *H. huberi* is assigned to Bartonian–Priabonian
6
7 (SBZ 17–20). This species is associated with *R. malatyaensis* and indicating the Bartonian age.
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14 Family Soritidae Ehrenberg, 1839

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16 Genus *Neorhipidionina* Hottinger, 2007

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18 *Neorhipidionina spiralis* Hottinger, 2007
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21 Figure 3G
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28 *Type species.* — *Rhipidionina williamsoni* Henson, 1948
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31 *Description.* — In the record species, the maximum diameter reaches 4.75 mm with a thickness
32
33 of 0.6 mm.
34

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36 *Remarks.* — The taxon found in this work is attributed to *Neorhipidionina spiralis* Hottinger
37
38 due to its more flattened in respect to *Neorhipidionina urensis* (Henson). The latter species is
39
40 showing an oval outline in cross section and its coils are more inflated and narrow. The
41
42 megalospheric form of *Neorhipidionina williamsoni* (Henson) is twice as large as the *N. spiralis*.
43
44 The recorded species is characterized by comparatively small size and long spiral early stage of
45
46 growth followed by rapidly flaring adult stage. According to Serra-Kiel *et al.* (2016), this species
47
48 ranges Lutetian in age. After Hottinger (2007), its biostratigraphic range extends from Bartonian
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50 to Priabonian. Based on faunal assemblage, the recorded species is assigned to the Bartonian
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52 stage.
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6 Family Soritidae Ehrenberg, 1839
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9 Genus *Archaias* de Montfort, 1808
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12 *Archaias operculiniformis* Henson, 1950
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15 Figures 3H1 and 3I
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21 *Type species.* — *Nautilus angulatus* Fichtel and Moll, 1798
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24 *Description.* — The equatorial diameter in megalospheric form reaches 3.2 mm in the record
25
26 sample.
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29 *Remarks.* — The *Archaias operculiniformis* Henson differs from *Archaias diyarbakirensis*
30
31 (Sirel) in its smaller size of test shell and proloculus. This species is characterized by fewer
32
33 pillars in respect of the *A. diyarbakirensis*. The basal layer ridges in the genus of *Penarchaias*
34
35 Hottinger could not be formed the pillars. According to Sirel (1998), Hottinger (2007), and Bassi
36
37 *et al.* (2007), the genus *Praearchaias* Sirel has the same structural elements as the *Archaias* De
38
39 Montfort. Hottinger (2007) mentioned the *Archaias operculiniformis* Henson from Bartonian-
40
41 Priabonian, but Serra-Kiel *et al.* (2016) extended this species to Rupelian in Oman. In this work,
42
43 the biostratigraphic range of recorded species can be attributed to the Bartonian stage.
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51 Family Soritidae Ehrenberg, 1839
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54 Genus *Penarchaias* Hottinger, 2007
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3 *Penarchaias glynnjonesi* (Henson, 1950)
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6 Figure 3H2
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12 *Type species.* — *Peneroplis glynnjonesi* Henson, 1950
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15 *Description.* — In the present species, the equatorial diameter of the test is 1.9-2.2 mm with a
16 thickness of 0.9 mm.
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20 *Remarks.* — The genus of *Penarchaias* (Henson) differs from *Archaias* De Montfort in the
21 absence of pillars. But in the point of view of morphological chambers and distribution of the
22 foramina, this genus is close to *Archaias* De Montfort. The *Penarchaias glynnjonesi* (Henson), is
23 recorded in the Bartonian stages in the studied area but its biostratigraphic range extends to
24 Oligocene (Hottinger, 2007; Serra-Kiel *et al.*, 2016).
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35 Class Globothalamea Pawlowski *et al.*, 2013
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37 Order Loftusiida Kaminski and Mikhalevich, 2004, *in*: Kaminski, 2004
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40 Family Coskinolinidae Moullade, 1965
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43 Genus *Coskinolina* Stache, 1875
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46 *Coskinolina perpera* Hottinger and Drobne, 1980
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50 Figures 4B and 4C
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3 *Type species.* — *Coskinolina liburnica* Stache, 1875
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6 *Description.* — In the record species, the axial length and basal diameter are 1.9- 2.0 mm and
7
8 1.7 -1.9 mm, respectively. The endoskeleton is only represented by irregular pillars and showing
9
10 a looser aspect. The pillars are discontinuous from one chamber to the next.
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12

13 *Remarks.* — Conical agglutinated shell shows pseudokeriothecal wall without beams and
14
15 intercalary beams. The endoskeleton is only represented by pillars and showing a looser aspect.
16
17 The wall is thicker than in all other species of the same genus. *Coskinolina perpera* Hottinger
18
19 and Drobne is ranged from Late Cuisian to Early Lutetian (Hottinger and Drobne, 1980).
20
21 Hottinger (2007) reported this species from Lutetian to Bartonian of Iran. According to Serra-
22
23 Kiel *et al.* (2016), the age of this species can be reached to Priabonian in Oman and Yemen. The
24
25 recorded species is assigned to the Bartonian (SBZ 17–18) in the studied area.
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33 Family Coskinolinidae Moullade, 1985
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36 Genus *Coskinolina* Stache, 1875
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39 *Coskinolina liburnica* Stache, 1875
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42 Figure 4A
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48 *Type species.* — *Coskinolina liburnica* Stache, 1875
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51 *Description.* — The axial length and basal diameter of the record sample are 2.1 mm and 1.9
52
53 mm, respectively. The pillars are densely arranged from one chamber to the next. There are not
54
55 exoskeletal elements (beams and rafters) in the marginal chamber.
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3 *Remarks.* — The identification of generic forms is distinguished by the lack of exoskeletal
4 structures in trochspiral nepionts. The *Coskinolina liburnica* Stache shows an acute cone angle.
5
6 The uniserial chambers are low and the pillars show a dense internal structure compared to all
7
8 other species of the same genus. According to Hottinger and Drobne (1980), *C. liburnica* is
9
10 assigned to the late Early Eocene (Cuisian) in the Island of Molat (Melada, *Croatian* Island in
11
12 the Adriatic Sea). Serra-Kiel *et al.* (1998) considered its biostratigraphic range from middle
13
14 Cuisian to Late Cuisian. Serra-Kiel *et al.* (2016) stated that the co-occurrence of *C. liburnica*
15
16 Stache with *Nummulites fabianii* (Prever) (SBZ 19) in the western Dhofar (Oman), indicates the
17
18 Periabonian age. As a result, the species of *C. liburnica* Stache can be extended from the late
19
20 Ypresian to the Priabonian. In east Iran, the *C. liburnica* Stache is reported from the Early
21
22 Eocene (Babazadeh, 2003; Schlagintweit and Hadi, 2018). In this work, this species is found in
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24 late Middle Eocene (Bartonian) sedimentary rocks.
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35 Family Orbitolinidae Martin 1890

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37 Genus *Barattolites* Vecchio and Hottinger, 2007

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39 *Barattolites trentinarenis* Vecchio and Hottinger, 2007

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42 Figures 4D and 4E
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49 *Type species.* — *Barattolites trentinarenis* Vecchio and Hottinger, 2007, figs. 12 a-b, 13 a-b, 14
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51 a-b, 15 a-m, 16 a-m, 17 a-r, 18 a-j
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3 *Remarks.* — The genus *Barattolites* Vecchio and Hottinger differs from the genus *Daviesiconus*
4 Hottinger and Drobne in having a long trochospiral early growth stage, the constant occurrence
5 of intercalary beams, and in the absence of marginal apertures. This genus differs from the genus
6 *Dictyoconus* Blanckenhorn and *Fallotella* Mangin due to the absence of horizontal partitions
7 (rafters) and to presence of a simple exoskeleton. It differs from *Coskinolina* Stache in having
8 simple radial partitions (beams and intercalary beams). The stratigraphic range of *Barattolites*
9 *trentinarensis* Vecchio and Hottinger covers the Ypresian to the Early Lutetian (Vecchio and
10 Hottinger, 2007). Based on foraminiferal association, the recorded species is assigned to
11 Bartonian in the studied area.
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28 Order Rotaliida Delage and Herouard, 1896

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30 Family Rotaliidae Ehrenberg 1839

31
32 Genus *Medocia* Parvati, 1971

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36 *Medocia blayensis* Parvati, 1971

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39 Figures 4F and 4G1
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45 *Type species.* — *Medocia blayensis* Parvati, 1971
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48 *Description.* — The diameter of the record sample is larger than 1 mm (ratio of equatorial to an
49 axial diameter about 2).
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53 *Remarks.* — The typical forms of *Medocia blayensis* Parvati are described from the Middle
54 Eocene (Lutetian - Bartonian) of France and Iran by Parvati (1971) and Hottinger (2007)
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3 respectively. The small forms ($d < 1\text{mm}$) of *Medocia blayensis* Parvati are reported by Vecchio
4 (2003) and Benedetti *et al.* (2011) from Ypresian of central and southern Italy. According to Le
5
6 Calvez and Blondeau (1978), *Medocia blayensis* Parvati together with *Alveolina elongate*
7
8 d'Orbigny, *Alveolina fusiformis* Stache and *Orbitolites cotentinensis* Lehmann found in Late
9
10 Lutetian sediments of the Atlantic Coast (Europe). Whereas, in the Trentinara Formation
11
12 (Southern Italy), this species is occurred in Late Early Eocene sediments. After Serra-Kiel *et al.*
13
14 (2016), the biostratigraphic range of *Medocia blayensis* Parvati is Early?-Middle Lutetian to
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16 Priabonian (SBZ 13?–SBZ 20). The biostratigraphic range of the recorded species is assigned to
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18 Bartonian in the studied area.
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28 Family Rotaliidae Ehrenberg, 1839

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30 Genus *Rotaliconus* Hottinger, 2007

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33 *Rotaliconus persicus* Hottinger, 2007
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36 Figure 4H
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42 *Type species.* — *Rotaliconus persicus* Hottinger, 2007
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45 *Remarks.* *Rotaliconus* Hottinger differs from all other calcarinids in its single interiomarginal
46 aperture and the presence of an enveloping canal system. It differs from rotaliids with umbilici
47 covered by funnel orifices, such as *Kathina* Smout and related forms, in having dorsal orifices of
48 the canal system and in the absence of continuity in the funnels of successive whorls (Hottinger,
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50 2007). In recoded specimens, the axial diameter of the test is larger than 1 mm (ratio of
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3 horizontal to an axial diameter about 1.4). The diameter of the base of the cone in the final whorl
4 measured 1.25 mm. According to Rahaghi (1980), its biostratigraphic range is Lutetian, but
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6 Hottinger (2007) extended its biostratigraphic range to Bartonian (SBZ 17-18). Serra-Kiel *et al.*
7
8 (2016) reported that the biostratigraphic range of this species could be extended from early?-
9
10 middle Lutetian to Priabonian (SBZ 13?–SBZ 20). The biostratigraphic range of the recorded
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12 species is considered to Bartonian in the studied area.
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20 **Discussion and Conclusion**

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26 The present work focused on the benthic foraminiferal biostratigraphy of the Jahrum Formation
27 in the Shahrekord region of southern Iran. According to Rahaghi (1978), the nummulite facies
28 with the *Assilina exponens* group and other nummulitids appear at the lower part of Jahrum
29 Formation during Early Eocene (Ypresian) to early Middle Eocene (Lutetian) in the type locality
30 of this formation in the Shiraz area. But the upper part of the Jahrum Formation consists of
31 porcelaneous foraminifera (miliolids), conical agglutinated foraminifera (coskinolinids), and rare
32 small rovalids. The miliolids along with small rovalids represent an oligotypic community of
33 epifaunal benthos and indicate a low energy depositional setting (Zamagni *et al.*, 2008). This
34 association occurs in near-shore lagoonal environments with muddy substrates and nutrient-
35 enriched conditions. The presence of *Macetadiscus* Hottinger, Serra- Kiel and Gallardo-Garcia
36 suggests relatively shallower water than *Alveolina* d'Orbigny. The *Macetadiscus* bearing facies
37 has a lateral paleoenvironmental relationship with the *Alveolina* facies. Apparently, the upper
38 part of the Jahrum Formation as the same as the Igualada basin (Northeastern Spain), is formed
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3 in a shallowing upward trend which represents a regressive cycle at the end of the late Middle
4 Eocene (Bartonian). The absence of larger benthic foraminifera such as *Nummulites* Lamarck
5 and *Alveolina* d'Orbigny in the study area can be matched with the different environments
6 because both taxa adapted to a reduced light condition in the open platform and occurred in
7 deeper water or offshore carbonate platform.
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17 In the point of view of paleontology, the *R. malatyaensis* is an index fossil for Bartonian
18 because it has a wide geographic distribution from western to central Neotethys (Hottinger,
19 2007). The co-occurrence of *R. malatyaensis* with *N. fabianii* (Perver) and *N. biedai* Schaub in
20 Arabil and Devely sections (Turkey) represents that the *Rhabdorites malatyaensis* (Sirel) is
21 ranged from Bartonian to Priabonian in age (Sirel, 2003). Afterward, some researchers, such as
22 Serra-Kiel *et al.* (2016) and Nafarieh *et al.* (2019), proposed that the biostratigraphic range of the
23 *R. malatyaensis* extends from Bartonian to Priabonian age. On the other hand, the occurrence of
24 *R. malatyaensis* seems to mark the level of *G. iranica* in the Shiraz area (west Iran). The range of
25 *R. malatyaensis* is shown to overlap that of *N. fabianii* (SBZ 19) (Serra-Kiel *et al.*, 1998).
26 According to Hottinger (2007), the type level of *G. iranica* is older than the range of *N. fabianii*
27 and maybe younger than the range of the *Assilina exponens* group (SBZ 13-17), which is
28 younger than SBZ 17. Also, the *R. malatyaensis* occurs in association with *M. vicensis*, in the
29 Bartonian sediments of the Igualada Basin of northern Spain (Romero *et al.*, 1999). Therefore,
30 the age of the *R. malatyaensis* is attributed to the Bartonian and this species can be considered as
31 an index fossil for Bartonian age. Thus, among the large benthic foraminifera in the study area,
32 only *R. malatyaensis* is known as a well-dated taxon. In the study area, the foraminiferal
33 association represents one concurrent range zone (Figure 5) which is compatible with the
34 biozone introduced by Hottinger (2007). Therefore the similar association of Bartonian benthic
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3 foraminifera has been compared with the established assemblage for the Shiraz area (south Iran),
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5 Dhofar section (Oman), and Socotra Island (Yemen) (Hottinger, 2007; Serra-Kiel *et al.*, 2016;
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7 Nafarieh *et al.*, 2019).
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23 agglutinated conical foraminifera (Coskinolinidae). The manuscript is benefited from the final
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27 appreciated.
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Figure captions

Figure 1. (A) Iran map showing the several sedimentary basin zones; (B) Location of the study area (Kuh-e-Soukhteh) in the High Zagros Zone on the Chaharmahal Bakhtiari Province (Zahedi and Rahmati Ilkhechi, 2006); (C) Location of the study area in Ardal geological map (1: 250 000).

Figure 2. Distribution of selected benthic foraminifera on stratigraphic section.

Figure 3. Photographs of foraminiferal species. (A) *Nurdanella boluensis* Özgen, equatorial section, As 14; (B) *Austrotrilina eocaenica* Hottinger, axial section, As 38; (C1) *Macetadiscus* cf. *incolumnatus* Hottinger, subaxial section and (C2) *Rhabdorites malatyaensis* (Sirel), equatorial section, As 38; (D) *Rhabdorites malatyaensis* (Sirel), axial section, As 43; (E1) *Praerhapydionina delicata* Henson, axial section, complete skeleton, As 50 and (E2) *Rhabdorites malatyaensis* (Sirel), axial section, As 50; (F) *Haymanella huberi* (Henson), axial section, As 55; (G) *Neorhipidionina spiralis* Hottinger, subaxial section, As 14; (H1) *Archaias operculiniformis* Henson, oblique section, (H2) *Penarchaias glynnjonesi* (Henson), axial section, As 77; (I) *Archaias operculiniformis* Henson, axial section, As 38; Bartonian, Scale bars: 1mm. The arrow shows the foraminiferal species in a thin section.

Figure 4. Photographs of foraminiferal species. (A) *Coskinolina liburnica* Stache, axial section, As 46; (B and C) *Coskinolina perpera* Hottinger and Drobne, B, Axial section, As 40, C, equatorial section, As 38; (D and E) *Barattolites trentinarensis* Vecchio and Hottinger, D, subaxial section, As 69, E, equatorial section; (F) *Medocia blayensis* Parvati, axial section, As 42; (G1) *Medocia blayensis* Parvati, subaxial section, As 41, (G2) *Macetadiscus* cf. *incolumnatus* Hottinger, subaxial section, As 41; (H) *Rotaliconus persicus* Hottinger, axial

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3 section, As 38, Bartonian, Scale bars: 1mm. The arrow shows the foraminiferal species in the
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5 thin section.
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8 Figure 5. Range chart of selected benthic foraminifera based on Romero *et al.* (1999); Sirel
9 (2003); Hottinger (2007); Serra-Kiel *et al.* (2016); Nafarieh *et al.* (2019) and study area. SBZ
10 zonation according to Serra-Kiel *et al.* (1998), the pale grey area shows concurrent range zone.
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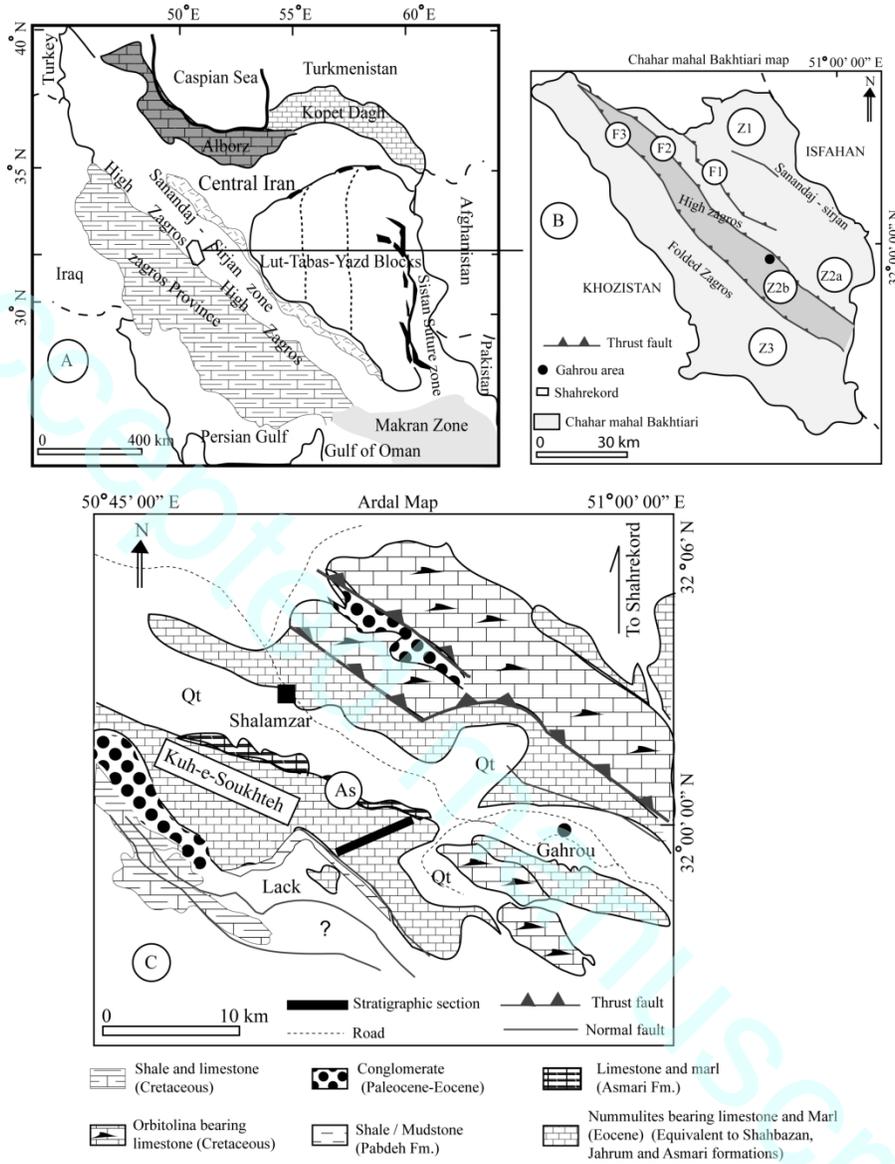


Fig. 1

Figure 1

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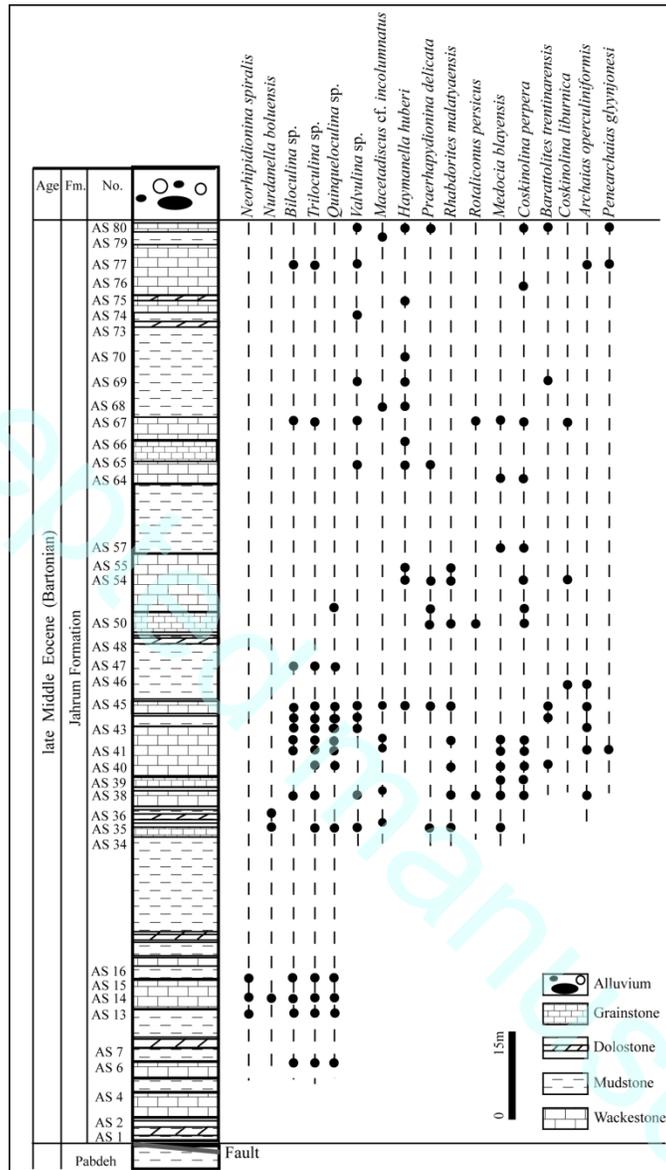


Fig. 2

Figure 2

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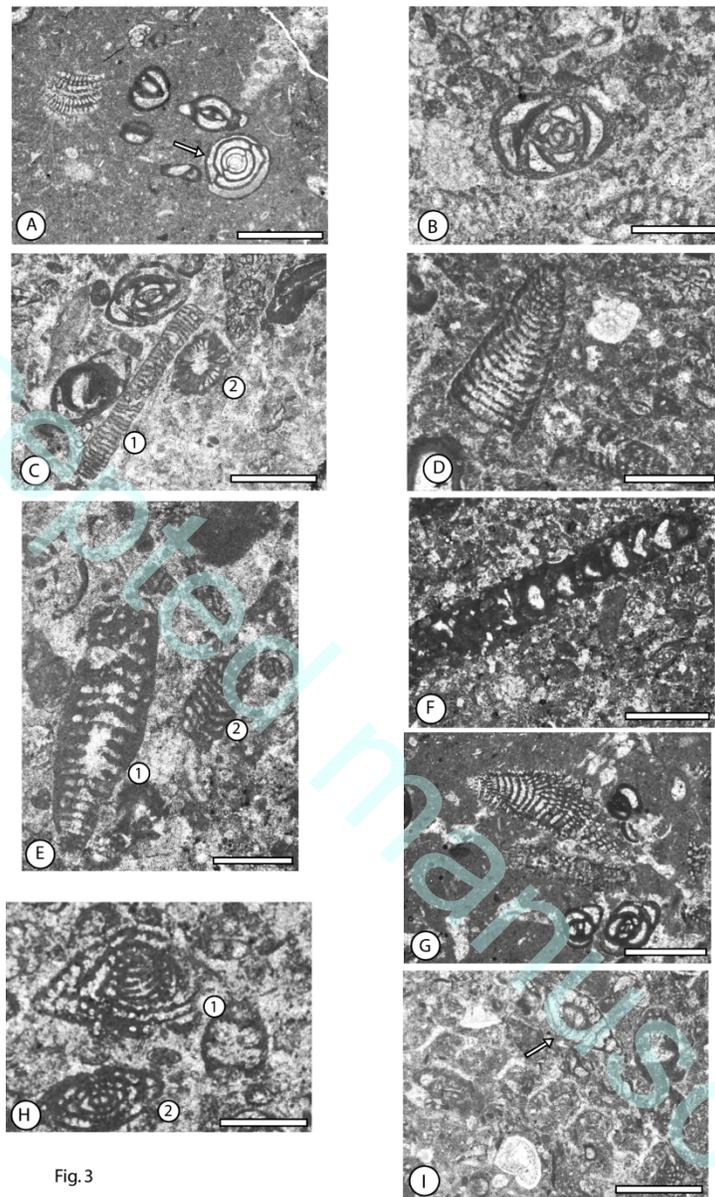


Fig.3

Figure 3

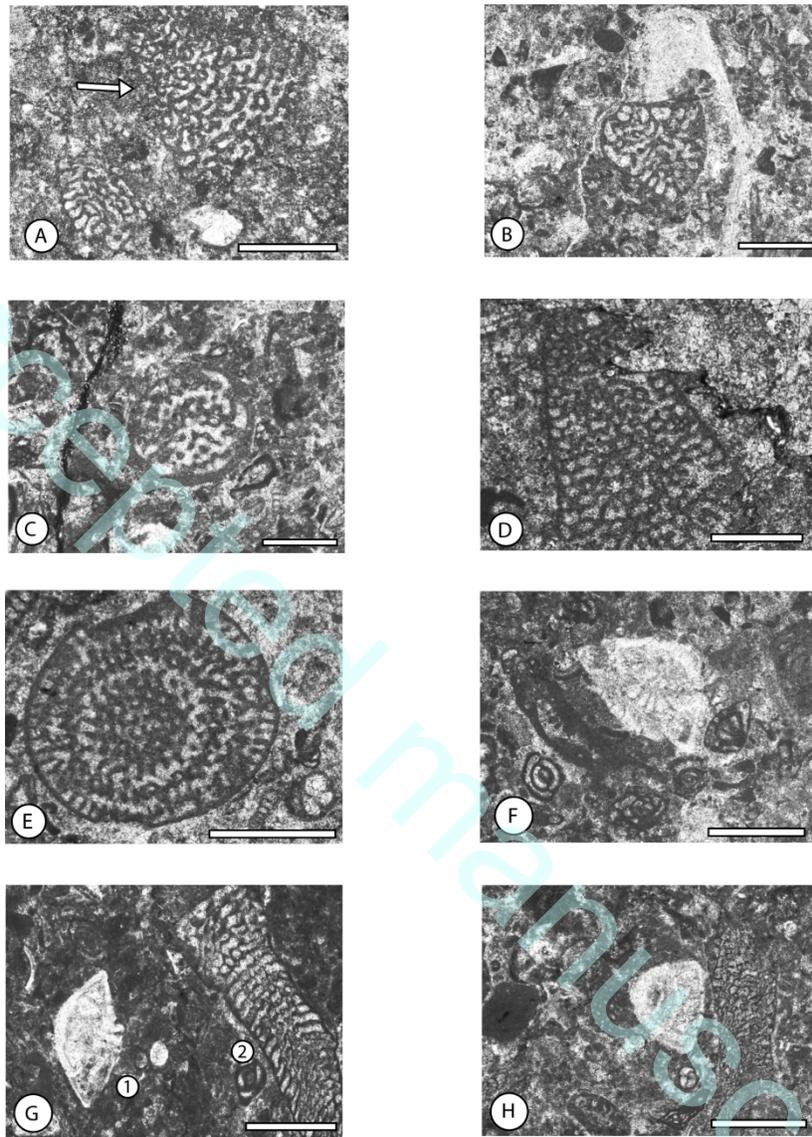


Fig.4

Figure 4

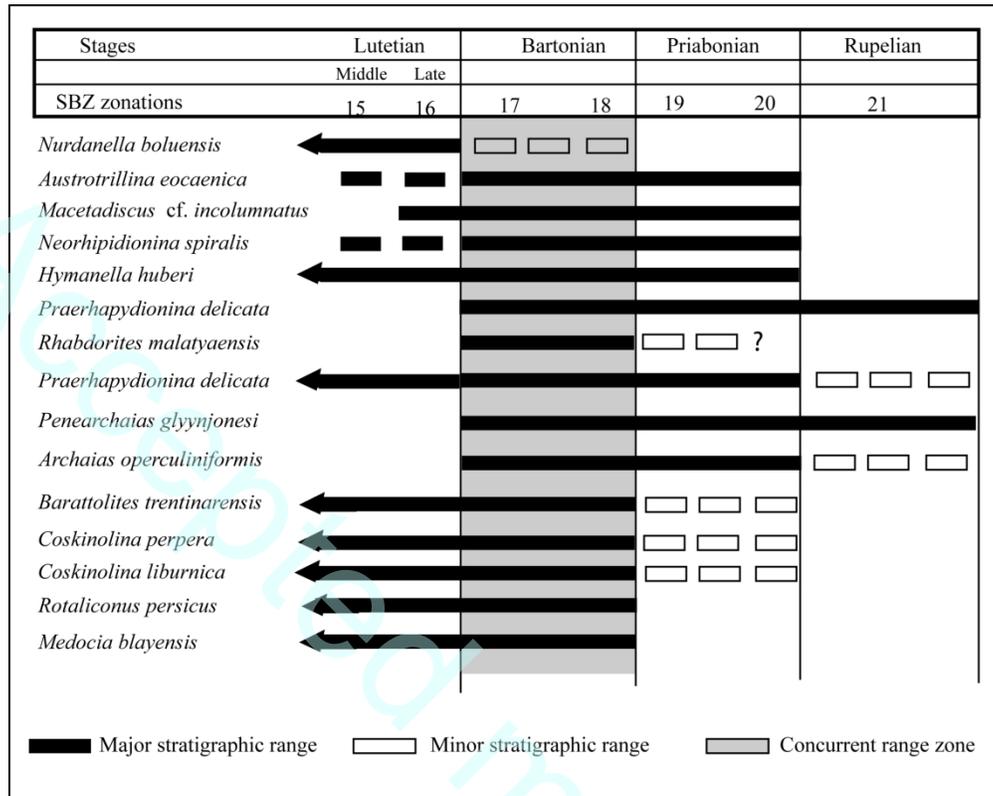


Fig. 5

Figure 5

196x165mm (300 x 300 DPI)