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The fossil on the cover is an adult example (T. TAKAHASHI coll.) of *Mikasaites orbicularis* MATSUMOTO (subfamily Marshallitinae, family Kossmaticeratidae) from the Lower Cenomanian (Cretaceous) of the Mikasa area, central Hokkaido. (photo by M. NODA, natural size)

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## 740. CRETACEOUS TETRAGONITID AMMONITE JAWS: A COMPARISON WITH MODERN NAUTILUS JAWS\*

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Abstract. The morphology and mineral composition of jaw apparatuses in Tetragonites glabrus, Gaudryceras denseplcatum, G. tenuiliratum and G. sp. are described. The studied material is well-preserved specimens in some calcareous nodules from the Turonian-Maastrichtian of Hokkaido and Sakhalin. The upper and lower jaws of these species consist of black outer lamella, reduced short inner lamella and covering rostra. The two lamellae are composed of carbonate apatite, and the rostra are calcareous covering. The lamellae and rostra may have been diagenetically replaced from chitin and aragonite respectively, judging from the comparison with modern cephalopod jaws. The calcareous rostra of the upper and lower jaws are analogous to rhyncholite (upper jaw) and conchorhynch (lower jaw) of modern Nautilus instead of coleoids in the internal microstructure, and mineral and chemical composition. These evidences suggest an analogy in feeding habits between tetragonitid ammonites and *Nautilus*, although the ammonites have larger and stronger jaws.

of

ammonites

#### Introduction

Anaptychi and aptychi are often preserved in ammonoid body chambers. They have been interpreted as ammonoid opercula (e.g. Retowski, 1891; Schmidt, 1928; Trauth, 1927–1936; Nagao, 1931a-c, 1932; Fischer & Fay, 1953; Arkell, 1957; Schindewolf, 1958) or jaw apparatuses (e.g. Meek & Hayden, 1864; Kennedy & Cobban, 1976). Recently, Closs (1967a, b) discovered several examples of anaptychus (upper jaw)—radula associations within the living chambers of an Upper Carboniferous ammonoid, Eoasianites (Glaphyrites) rionegrensis Closs, and Lehmann (1967, 1971b, 1979) reported in situ aptychi or anaptychi and radulae in some Jurassic ammonites. Thus, aptychi and anaptychi have been confirmed as the lower jaws

to paratuses in modern coleoids were studied

Gasiorowski, 1973; Shimansky, Müller, 1974; etc.) or coleoid jaws (Biguet, 1819; Gaillardot, 1824; de Blainville, 1827; etc.), but some were considered to belong ammonoids (Schmidt-Effing, Mundlos, 1973; Tanabe et al., 1980; etc.). Almost all cephalopods have jaw ap-The taxonomic paratuses and radulae. and ecological implications of jaw ap-

Jurassic

examples

Recently

On the

1962:

1972:

Tetragonites.

in

Kaiser & Lehmann, 1971).

apparatuses of Lytoceratacea.

paratuses of Cretaceous

(Lehmann, 1967, 1970, 1971a, 1972, 1976,

1978, 1979; Lehmann & Weitschat, 1973;

Zakharov (1979) described the jaw ap-

Lehmann et al. (1980) described the jaw

other hand, rhyncholite and conchorhynch

were interpreted as calcified elements of

nautiloid jaws (d'Orbigny, 1825; Foord,

1891; Rutte, 1962; Teichert et al., 1964;

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by Clarke (1962), Nixon (1969) and others, and those in *Nautilus* by Okutani & Mikami (1977) and Saunders *et al.* (1978).

Most previous works on ammonoid jaws have been done mainly on the basis of the Jurassic materials from Europe. However, Nagao (1931a-c, 1932) reported Cretaceous aptychi and anaptychi from Hokkaido and south Sakhalin, and Trauth (1935) redescribed Nagao's *Neoanaptychus*.

Kanie *et al.* (1978) preliminarily reported the jaw apparatuses of some Late Cretaceous ammonites from Hokkaido and Sakhalin. They recognized in some specimens that a pair of lower and upper jaws are preserved in the body chamber; calcareous coverings similar to rhyncholite and conchorhynch exist in the rostra of



Text-fig. 1. Locations of the specimens studied and exposures of Cretaceous deposits (solid area). N: Naibuchi, K: Tombetsu, R: Obira, As: Ashibetsu, Ik: Ikushumbetsu, Y: Oyubari, Kd: Furenai, and U: Urakawa areas.

lower and upper jaws; the ammonite jaws are biomineralogically comparable with those of extant *Nautilus*. Furthermore, the tetragonitid jaws were interpreted to be analogous to those of *Nautilus* owing to similar feeding habit (Kanie & Tanabe, 1979; Tanabe *et al.*, 1980).

This paper presents a description o. the general morphology of the teragonitid ammonite jaws from the Upper Cretaceous of Hokkaido and Sakhalin (Text-fig. 1), and compares them with the jaw apparatuses of modern *Nautilus* by some mineralogical analysis. On the basis of the results obtained, the tetragonitid jaws are restored.

The depositories of the specimens examined are as follows (abbreviations are shown in parentheses):

- Yokosuka City Museum, Yokosuka 238 (YCM) Department of Geology, Kyushu University, Fukuoka 812 (GK)
- Department of Geology and Mineralogy, Hokkaido University, Sapporo 060 (UH)
- Japanese Expert Consultation on Living Nautilus. c/o Yokosuka City Museum (JECOLN)
- Y. Kawashita's private collection, 2 A1-5, Tomatsu-Chiyoda-cho, Mikasa 068-22
- M. Nihongi's private collection, 8S, 22W, Chuo, Sapporo 064 (MN)
- T. Takahashi's private collection, 28-109, Hanazono-cho, Mikasa 068-22

#### General morphology

#### Comparison with modern cephalopods

Identification of lower and upper jaws.— The lower jaw is distinguished from the upper one by several characters (Textfigs. 2, 3), namely, the larger size of lamellae (compared with the height and width of living chambers), the broader lamella angle (angle between left and right sides of the outer lamellae) and less



Text-fig. 2. Various kinds of basic morphologies in modern coleoid jaws. Upper jaw of the Vampyromorpha cited from Iverson & Pinkas (1971) and lower jaw from Clarke (1980). Octopoda, Teuthoidea and Sepioidea cited from Tanabe *et al.* (1980). Terminologies are based on Clarke (1962). w: wing, c: crest, lw: lateral wall, h: hood.

extended hood area (Clarke, 1962 for coleoids; Okutani & Mikami, 1977 and Saunders *et al.*, 1978 for *Nautilus*).

Morphological comparison with modern cephalopods.-The morphological features of the tetragonitid jaws (Text-figs. 4, 7) can be correlated with those of living Nautilus rather than with those of modern coleoides. These ammonite jaws are characterized by the presence of rhyncholites and conchorhynchs at the anterior part of the jaws (Tanabe et al., 1980), The Nautilus jaw is also covered with rhyncholite and conchorhynch (Text-fig. 3; Okutani & Mikami, 1977; Saunders et al., 1978). Those of coleoids, however, lack calcareous coverings visible with the naked eye (Text-fig. 2). The lower jaw

of the Tetragonitidae is morphologically similar to those of *Nautilus* and *Vampyroteuthis infernalis* in the scarcely extended crest and lateral wall. In the Tetragonitidae the upper jaw more or less resembles the lower jaw except for the more extended wings covering the reduced short lateral walls and crest, whereas in these living cephalopods the upper and lower jaws are morphologically quite different from each other.

#### Description of jaw apparatuses

The jaw specimens examined are attributed to *Gaudryceras* and *Tetragonites* (Tetragonitidae : Lytoceratina). They were previously named *Neoanaptychus* by Nagao



Text-fig. 3. Basic morphology of modern *Nautilus* jaw. L: length of lateral wall, W: length of wing, Hd: length of hood, Hi: height, B: breadth,  $\alpha$ : wing (lateral wall) angle,  $\beta$ : jaw angle. Terminologies are mostly based on Okutani & Mikami (1977) and Saunders *et al.* (1978). Solid area is covered with a calcareous covering.

(1931a) and Anaptychus by Trauth (1935).

Family Tetragonitidae Hyatt, 1900 Subfamily Gaudryceratinae Spath, 1927 Genus *Gaudryceras* Grossouvre, 1894

#### Jaws of Gaudryceras spp.

*General characters*: The lower and upper jaws are of *Anaptychus* type. They are broader than high. Lower jaw is larger and broader than upper one. Both



Text-fig. 4. Restoration of *Gaudryceras* jaw and the measurements. W: length of wing, Hi: height, B: breadth, Hd: length of hood,  $\alpha$ : wing angle,  $\beta$ : jaw angle. Solid area is covered by a calcareous covering. U1 and L3 (Lehmann *et al.*, 1980).

lower and upper jaws are constructed by wing, hood and rostrum with calcareous covering and crest plus minor lateral wall beneath them. Wing angle ( $\alpha$ , angle between left and right sides of the wings; Text-fig. 7) is wide in the lower jaw. Jaw angle  $(\beta)$  is obtuse; the curve between rostrum and wing margin is gentle without notch. The black lamellae are made up of apatitic mineral and covered with a thin calcite layer.

*Remarks*: The specific distinction of these jaws in accordance with the classification of *Gaudryceras* is difficult at present.

Upper jaw (Pl. 39, Figs. 1-3,5; Text-fig. 4-U)

Previously described material: The following specimens are attributable to the upper jaws of Gaudryceras spp.

- 1931c. Neoanaptychus sp. form 7, Nagao, p. 216, pl. 15, fig. 6.
- 1935. Anaptychus tenuiliratus (Nagao), Trauth, pp. 452-454.

Material: Attributed to Gaudryceras denseplicatum (Jimbo) in situ. GK. H8087. Loc. R4041p2, the upper part of the Obirashibe River, Obira area, northwestern Hokkaido. Lower Santonian (H. Hirano & K. Tanabe coll.); to G. tenuiliratum Yabe in situ. UH3191. Loc. the Ikushumbetsu River about 200 m east from the junction with the Kikumezawa, Ikushumbetsu area, central Hokkaido. Santonian (R. Saito coll.); to ? G. tenuiliratum Yabe. GK. H2262. Loc. N146, the lower part of the Santan, Naibuchi area, south Sakhalin. Santonian (T. Matsumoto coll.); to ? G. tenuiliratum Yabe. Loc. Kawakami, the Obirashibe River, Obira area. Sahtonian ? (T. Takahashi coll.); to ? G. sp. Loc. Y5251a, the lower part of Hakkinzawa, Oyubari area, central the Hokkaido. Turonian (T. Matsumoto, H. Hirano & K. Tanabe coll.); to ? G. sp. Loc. R4040b. the upper part of the Obirashibe River, Obira area, northwestern Hokkaido. Upper Coniacian (H. Hirano & K. Tanabe coll.); to ? G. sp. MN6, 83c. Loc. the upper part of the Yubarigoenosawa, Ikushumbetsu area. Coniacian ? (T. Nihongi coll.); to ? G. sp. (=Neoanaptychus form 7) UH3193. Loc. Kikumezawa, Ikushumbetsu area. Santonian ? (R. Saito coll.); to ? G. sp. GK. H3220. Loc. Ur512Y4, the Ikandai Valley, Urakawa area, southerncentral Hokkaido. Lower Campanian (T. Matsumoto coll.); to ? G. sp. GK. H8088. Loc. K57, Heitarozawa, Tombetsu area, northern Hokkaido. Upper Maastrichtian (T. Matsumoto, H. Okada, I. Orita & T. Muramoto coll.).

*Description*: The wings are triangular and make up the greater part of the upper jaw. Wing surface is marked by concentric costae and numerous growth lines of variable prominence. The wings are slightly compressed towards the flank of the living chamber, and combine a hood at the antero-dorsal part, being narrowly flat at the dorsal part; growth lines gently concave toward anterior direction. In some specimens (e.g., Pl. 39, Fig. 1b), folds or a groove are found; wings posteriorly projected in later growth stage. Crest and lateral wall are small. Wing

Specimen	W	Hd	Hi	В	α	β	Lh	(Hi/LH)
UH3193	56.9	38.4	37.0	58.0	60°	165°		
MN6	58.8+	41.5	29.2	40.0	49°	155°		_
GK.H8086	34.7	31.2	25.7	30.9*	51°	170°		—
R4040b	33.4 +	30.7	27.4 +	27.0*	65°	165°	_	
GK.H3220	28.7	24.3	25.0 +	32.9	48°	165°		
GK.H8087	23.7	18.7	23.1 +			135°	29.2	(0.64)
GK.H2262	16.5	14.9	18.8	14.0	47°	160°	_	_
UH3191	12.7	10.9	11.6	14.4 +	87°*	170°	18.9	(0.61)

Table 1. Measurements (in mm) of upper jaws of Gaudryceras spp.

W: wing, Hd: hood, Hi: height, B: breadth,  $\alpha$ : wing angle,  $\beta$ : jaw angle, Lh: height of living chamber, \*: deformed specimen.

angle is 47-65°.

The rostrum projects toward the ventral side and is covered with a calcite layer. In specimen MN6 (Pl. 39, Fig. 1), the poorly preserved rostral tip fills inside of the rostrum. The rostrum, originally sharp-topped, decreases in thickness toward the wings; the outer surface of the wings is covered with a light-colored, thin (approximately 500  $\mu$ m thick) calcite layer. The jaw angle is generally obtuse (135-170°), without a notch (Table. 1).

*Remarks*: In comparison with *Gaudry-ceras*, the upper jaw of *Nautilus* has less extended wings, larger lateral wall and narrower jaw angle (118-119°). The upper jaw of *Tetragonites* is clearly distinguished from that of *Gaudryceras* in the higher, narrower and more rounded outline.

Lower jaw (Pl. 39, Figs. 4-6; Text-figs. 4-1, 5, 6-1)

Previously described material: The following specimens are attributable to the lower jaws of Gaudryceras spp.

1931a. Neoanaptychus tenuiliratum Nagao, pp. 106-109, fig. 1.

- 1931c. Neoanaptychus tenuiliratus Nagao, p. 216, pl. 15, fig. 1.
- 1935. Anaptychus tenuiliratus (Nagao), Trauth, pp. 452-454.

Material: Attributed to Gaudryceras denseplicatum (Jimbo). GK. H8054. Loc. R2101p, the upper part of the Obirashibe River, Obira area, northwestern Hokkaido. Lower Turonian (T. Matsumoto & H. Hirano coll.); to G. tenuiliratum Yabe. GK. H8049. Loc. R2251b, the middle part of the Kamikinenbetsu River, Obira area. Upper Turonian (T. Matsumoto & H. Hirano coll.); to G. sp. GK. H8059. Loc. Y1664a, the upper part of the Shiyubari River, Oyubari area, central Hokkaido. Santonian (Y. Haraguchi coll.); to G. sp. GK. H8048. Loc. Kd905, Furenai area, central Hokkaido. Campanian (T. Matsumoto coll.); to ? G. denseplicatum (Jimbo). YCM. Ur307002. Loc. U307p, the Ikandai Valley, Urakawa area, southern-central Hokkaido. Lower Campanian (Y. Kanie coll.); to ? G. sp. GK. H8045. Loc. As1140, the lower part of the Souashibetsu River, Ashibetsu area, central Hokkaido. Coniacian (T. Matsumoto coll.).

Description: Wings are triangular and make up the greater part of the lower jaw. Wing surface is marked by concentric costae and numerous growth lines. Wings are slightly compressed toward the side of living chamber; the hood combines wings at the antero-ventral part, being flat, occasionally folded at the ventral part (Text-fig. 4-L2); growth lines gently concave toward the aperture. Wing angle is 80-87°. Wings are made up of carbonate apatite in the specimen YCM. Ur307002.

Lateral wall is small. Thick calcareous substance (rhyncholite) covers entire lateral wall (Text-fig. 5). Lateral wall continues to rostrum at the tip of beak. The rostrum is covered with a calcareous layer forming a beak which consists of an alveolus structure of 2 mm wide at the dorsal part, though decreasing in thickness to 0.5 mm toward posterior part of the wing and hood areas (Text-fig. 6-1).

*Remarks*: The general morphology of the lower jaw in *Gaudryceras* is similar to that of *Nautilus* except for the longer wing, partly flattened hood, and larger

Table 2. Measurements (in mm) of lower jaw of Gaudrycers denseplicatum (?).

Specimen	W	Hd	Hi	В	α	β	Lh	(Hi/Lh)
YCM. Ur 307002	37.8	25.0+	34.0*	51.0+	80°	140°		_

Symbols as in Table 1. \* deformed specimen.



Text-fig. 5. Crest and lateral wall structure (arrow) of the lower jaw probably attributable to *Gaudryceras* sp., GK. H8045, from loc. As1140, the lower part of the Souashibetsu River, Ashibetsu area. Coniacian.  $\times 2.3$ . Photo by K. Tanabe.



Text-fig. 6. Alveolus structure of Gaudryceras denseplicatum (1), YCM. Ur307002, from loc. U307p, the Ikandai Valley, Urakawa area. Lower Campanian (Tanabe et al., 1980) and Nautilus macromphalus (2), JECOLN.M5, from Nouméa, New Caledonia. Both are the lower jaws (dorsal views). Photos by Kanie.

jaw angle. The lower jaw of *Tetragonites* is clearly distinguished from that of *Gaudryceras* in the greater height, narrower wing angle, and larger jaw angle.

#### Subfamily Tetragonitinae Hyatt, 1900 Genus *Tetragonites* Kossmat, 1895

#### Jaws of *Tetragonites glabrus* (Jimbo, 1894)

General characters: Lower jaw is

broader than high. It is slightly larger than upper jaw. The breadth of upper jaw is approximately equal to or smaller than height. The jaw plate consists of outer lamellae (wing and hood) and inner lamellae (minor crest plus lateral wall). Wing angle ( $\alpha$ ) is wide in lower jaw. Jaw angle ( $\beta$ ) is obtuse to rectangular; the curve between rostrum and wing margin is gentle without notch. The wing is mainly composed of apatitic mineral with a thin covering of calcite layer.

Upper jaw (Pl. 40, Figs. 1-3; Text-fig. 7-U)

Material: Attributed to Tetragonites glabrus (Jimbo) in situ. GK. H2073. Loc. N143a, the lower part of the Santan, Naibuchi area, south Sakhalin. Lower Santonian (T. Matsumoto coll.); to T. glabrus (Jimbo) in situ. GK423. Loc. Sakhalin. Upper Cretaceous (collector unknown); to ? T. sp. Loc. the lower part of the Kaneobetsu, Oyubari area, central Hokkaido. Turonian (Y. Kawashita coll.).

*Description*: The wings are triangular and make up the greater part of the upper jaw. Wing surface is marked by concentric costae and numerous growth lines. Wings are slightly compressed toward the side of living chamber; the hood is flat or waved in dorsal part; growth lines concave toward the anterior part. Crest and lateral wall unclear. Wing angle is approximately 55°.

The rostrum is sharply projected toward the ventral side and covered with a thick calcareous layer (rhyncholite), decreasing the thickness toward the wing; outer surface of wings is covered with a thin (ca. 500  $\mu$ m thick) calcite layer (Pl. 40, Fig. 2). The jaw angle is approximately 55° in Kawashita's specimen, without a notch (Table 3).

*Remarks*: The lateral wall and crest of *Tetragonites* are reduced to a short



Text-fig. 7. Restoration of *Tetragonites* jaw and the measurements. Legends are shown in text-figure 4. U1 and L3 (Lehmann *et al.*, 1980).

inner lamella as compared with those of the upper jaw of *Nautilus*.

Lower jaw (Pl. 40, Figs. 3-5; Text-fig. 7-L)

Previously described material: The follow-

ing specimens are attributable to the lower jaws of *Tetragonites glabrus*.

- 1931a. Neoanaptychus tenuiliratum Nagao, pp. 106-109, fig. 2.
- 1931c. Neoanaptychus sp. form 5, Nagao, p. 216, pl. 15, fig. 4.

Specimen	W	Hd	В	Hi	α	β	Lh	(Hi/Lh)
Kawashita' coll.	<sup>'s</sup> 32.3	27.9	20.0*	28.0	55°	100°	_	_
GK423	17.3 +	13.8+	—	—	·	-	38**	(0.48)

Table 3. Measurements (in mm) of upper jaw of Tetragonites glabrus.

Symbols as in Table 1. \* deformed specimen; \*\* restored.

Table 4. Measurements (in mm) of lower jaws of Tetragonites glabrus.

Specimen	W	Hd	В	Hi	α	β	Lh	(Hi/Lh)
GK.H8043	28.9	23.2	18.9*	23.0*		135°		_
GK.H8049	26.2		20.0+	30.9*	85°	115°		
UH3194	24.6	19.3	13.4	30.5+	95°	100°		
MN83a	16.6	10.4	11.0*	27.0 +	95°	150°	—	—
GK.H2073	10.0	10.3	7.2*	25.0*	70°		37**	(0.36*)
UMUT.MM 6999	14.6	13.4	7.8	21.0	90°	110°	_	

Symbols as in Table 1. \* deformed specimen; \*\* restored.

- 1931c. *Neoanaptychus* sp. form 1, Nagao, p. 216, pl. 15, fig. 2.
- 1931c. Neoanaptychus sp. form 4, Nagao, p. 216, pl. 15, fig. 3.
- 1935. Anaptychus tenuiliratus (Nagao), Trauth, pp. 452-454, pl. 14, fig. 6.
- 1935. Anaptychus tenuiliratus var. lata Trauth, pp. 455-456, pl. 14, figs. 8, 9.
- 1979. Anaptychus sakhalinensis Yu. Zakharov, p. 87, pl. 20, fig. 2.

Material: Attributed to Tetragonites glabrus (Jimbo). GK. H8049. Loc. R2251b, the middle part of the Kamikinenbetsu River, Obira area, northwestern Hokkaido. Upper Turonian (H. Hirano & K. Tanabe coll.); to T. glabrus (Jimbo) in situ. GK. H2073. Loc. N143a, the lower part of the Santan, Naibuchi area, south Sakhalin. Santonian (T. Matsumoto coll.); T. glabrus (Jimbo) in situ. GK. H8043. Loc. N382, the Juhachirinpan-Ichinosawa, Naibuchi area. Santonian (T. Matsumoto coll.); to ? T. sp. MN83a. Loc. Takishita-Nakakinenbetsuzawa, the middle part of the Obirashibe River, Obira area. Upper Turonian (M. Nihongi coll.); to ? T. sp. (=Neoanaptychus form 5)UH3194. Loc. Kikumezawa, just above the junction with the Ikushumbetsu River, Ikushumbetsu area, central Hokkaido. Santonian (R. Saito coll.); to ? T. sp. UMUT.MM6999. Loc. Na & H-Cr738, the second eastern tributary of the Naibuchi River, Naibuchi area. Santonian (M. Kawada coll.).

Description: The wings are quadrate to triangular and make up the greater part of the lower jaw. Wing surface is marked by concentric costae and numerous Wings are gently comgrowth lines. pressed toward the side of living chamber: the hood is flat and sometimes folded at the ventral part (Nagao, 1931c, pl. 15, fig. 4); growth lines gently concave toward the anterior part. Wing angle is 90-95° (Table 4). Crest and lateral wall are small, existing as inner lamella. Thin calcite layer covers the greater part of outer surface of the wings. Wings are continuous to the rostrum at the tip of the beak (Pl. 40, Fig. 5c). The rostrum is sharp-topped, but denticles as cutting edge are not preserved; rostrum is covered with a calcite layer, the thickness of which decreases toward the posterior part of wings.

	Specimen		W (L)	Hd	В	Hi	α	β	Lh	(Hi/Lh)
JECOLN. M4	up.	(31.3)	19.0	17.0	25.3	33°	119°	04	(0.27)	
	M4	low.	28.9	19.5	33.0	24.5	52°	102°	94	(0.26)
	JECOLN.	up.	(34.0)	21.5	27.2	25.4	31°	118°	00	(0.27)
M5	low.	29.0	18.3	30.6	23.7	51°	102°	93	(0.26)	

Table 5. Measurements (in mm) of jaw appoffaratuses of adult Nautilus macromphalus.The specimens are from of Nouméa, New Caledonia.

W: length of wing (lower jaw), L: length of lateral wall (upper jaw), Hd: length of hood, B: breadth, Hi: height,  $\alpha$ : wing (lower jaw) or lateral wall (upper jaw) angle,  $\beta$ : jaw angle, Lh: height of living chamber, as illustrated in Text-figure 3.

Table 6. Comparison of jaw size of tetragonitid ammonites and modern Nautilus.

Genus		Upper	Lower
Gaudryceras	Hi/Lh W/Lh	0. 64 61 (2. 4-2. 3) 0. 81 67 (2. 3-2. 0)	
Tetragonites	Hi/Lh W/Lh	$\begin{array}{c} 0.48(1.8) \\ 0.27(0.8) \end{array}$	0.47 (1.5-1.3) 0.19*()
Nautilus	Hi/Lh WL/Lh	0.27(1.0) 0.3633(1.0)	$\begin{array}{c} 0.26 & (1.0) \\ 0.3631 & (1.0) \end{array}$

Hi: height, W: length of wing, L: length of lateral wall, Lh: length of living chamber. The number in parentheses is the ratio with that of *Nautilus*. \*: deformed specimen.

*Remarks*: The general outline of the lower jaw of *Tetragonites* is closely similar to that of *Nautilus* except for the narrower wing angle (51-52°) and smaller size (Table 5).

*Discussion*: The upper jaw plates of tetragonitid ammonites are morphologically more similar to those of octopods than decapods, and they resemble those of *Nautilus* more closely than octopods (Text-figs. 2, 3). Namely, the crest and lateral wall of the lower jaw of *Nautilus* are shortly reduced as in tetragonitid ammonites. Meanwhile, the upper jaw of modern *Nautilus* consists of wide inner lamellae (lateral wall and crest) and shortly reduced outer lamellae (hood and wing).

The outer and inner lamellae are morphologically similar between the upper and lower jaws in tetragonitid ammonites. The ratio of jaw height to living chamber height is less than 100 percent; for example, the upper jaw of *Gaudryceras* is 61-64%; the upper jaw of *Tetragonites* is 48% and lower jow 47%; the upper jaw of *Nautilus macromphalus* is 27% and lower jaw 26% (see Tables 5, 6).

Table 6 gives the size of the jaw apparatuses of tetragonitid ammonites in relation to the living chambers. The size of the jaw apparatuses in *Gaudryceras* is 1.7-2.4 times and in *Tetragonites* is 1.3-1.8 times as large as that of adult *Nautilus macromphalus*.

Sharp denticles are present on the alveolus structure in living *Nautilus*, but unknown in the Tetragonitidae. However, it may be due to the matter of fossilization (Text-fig. 6-2).

#### Mineral composition

The mineral composition of the jaw plates of tetragonitid ammonites was



Text-fig. 8. X-ray powder diffraction patterns of the lower jaw attributed to *Gaudryceras denseplicatum*. 1. Carbonate apatite, 2. calcite. The specimen examined is preserved in a calcareous nodule, from loc. U307p, Ikandai Valley, Urakawa area, southern-central Hokkaido. All pattens are obtained by Cu/Ni radiation. Chart speed 20 mm/min. Scanning speed 2°/min. Analyzed by A. Kato and S. Matsubara.

analyzed by an X-ray powder diffraction method. The patterns are compared with those of modern coleoids by the same The specimens examined are method. the lower jaws of Gaudryceras sp., Tetragonites glabrus, Nautilus macromphalus and Sepia sp. It was found that the calcareous covering of the rostrum in Gaudryceras is made up of calcite (Textfig. 8-2), whereas the black outer lamella (wing) is composed mainly of well-crystallized carbonate apatite (Text-fig. 8-1). The whole surface of the outer lamellae in Gaudryceras (Pl. 39, Fig. 1b) and Tetragonites (Pl. 40, Fig. 2) is covered with a calcite layer about 500 µm thick. The layers grades into the calcareous covering of the rostrum. The lower jaw plate of Nautilus consists of outer and inner lamellae, calcareous rostrum, and denticles. The former two are made up of chitin (Saunders et al., 1978), but the X-ray powder diffraction studies show that almost all materials display broad bands near d=4.4 Å only as represented by Text-figure 9-2, also that of Sepia (Text-



Text-fig. 9. X-ray powder diffraction patterns of the lower jaws attributed to two kinds of modern cephalopods. 1, 2. Nearly amorphous materials with broad bands about d=4.4 Å only, 3. calcite, 4. aragonite, 5. aragonite and calcite (heated by Y. Fukuda). Sepia sp. from Mexico Bay. Nautilus macromphalus from Nouméa, New Caledonia. Patterns were taken under the same recording conditions as those in Text-figure 8.

fig. 9-1). The calcareous covering of the rostrum is composed of aragonite (Text-fig. 9-4), but the denticles are of calcite (Text-fig. 9-3). Another test showed that a pair of the aragonitic rostra changed into calcite by heating at  $300^{\circ}$ C in half an hour (Text-fig. 9-5). The wing surface of the jaw of *Nautilus macromphalus* shows similarities with that of *Tetra-gonites*, a prismatic layer covers the outer surface of the wings of *Nautilus macrom* 

phalus (Fukuda, 1978).

Discussion: The wing surface of the upper and lower jaws of tetragonitid ammonites are covered with a calcite layer, approximately 500  $\mu m$  in thickness. Hölder (1958) reported that a thin calcareous layer covers the outer surface of a Jurassic anaptychus, probably attributable to Psiloceras torus d'Orbigny. Recently, Zakharov (1979) described the calcareous covering, 100-200  $\mu$ m thick, on the lamella of the lower jaw in Tetragonites popetensis Yabe. The outer lamellar surface (wing) of the lower jaw of modern Nautilus is underlain by an aragonite layer, 50-75  $\mu$ m thick (Fukuda, personal comm.). Its thickness is about 1/10 of that of tetragonitids. The aragonite layer covers the inner side of broadly extended lamellae. Therefore, the upper jaw of modern Nautilus is mostly occupied by the inner lamellae (crest and lateral wall). The outer lamellae of the upper jaw in modern cephalopods are generally reduced in length. Both the upper and lower jaws of tetragonitid ammonites are characterized by similar morphology. The calcified denticles of the tetragonitid jaws were originally made of calcite as in Nautilus, the rostrum and the outer surface of wings consisted of aragonite, and the wings were made up of chitin. During the course of fossilization, the rostrum and outer surface may have changed into calcite and the wings into carbonate apatite. This suggests that aragonite has changed into calcite during diagenesis. Such a conversion of aragonite into calcite is also indicated in the septa of Subptychoceras sp. (specimen YCM. Ur307001) preserved in the same nodule.

#### Restoration of jaws

The restoration of tetragonitid ammonite jaws is undertaken on the basis of the above mentioned observations and some anatomical studies of living *Nautilus*.

Saunders et al. (1978) discussed the function of the calcareous coverings of the rostral part, rhyncholite (upper jaw) and conchorhynch (lower jaw). in modern Nautilus and the extinct Nautilida. Tanabe et al. (1980) clarified some of the previously known fossil rhyncholites and conchorhynchs belong to such Cretaceous ammonites as Gaudryceras, Tetragonites, Zelandites and Neophylloceras. Morphologically the tetragonitid ammonite jaws are most similar to those of living Nau-However, the analogous internal tilus. structure is observed between upper and lower jaws of the ammonites in contrast to the great difference in Nautilus. The outer lamella (wings) of the upper jaws of Nautilus is reduced in length and the inner lamella (lateral wall) extends broadly. The lower jaw of modern coleoids is composed of lateral wall instead of wing.

Text-fig. 10 shows the articulation system of the upper and lower jaws of Nautilus pompilius in a buccal mass in the anterior part of the soft body. The calcareous covering, namely rhyncholite (rh) and conchorhynch (co), is shown as solid on soft X-ray photographs (Text-fig. 10), whereas the lamellae mostly made of chitin are faintly shadowed. Text-fig. 10-2 shows the articulated upper and lower jaws of Nautilus macromphalus. The muscle system between the upper and lower jaws of the present two species of Nautilus is designed in Text-figure 11 on the basis of data above mentioned and some anatomical studies. Namely, the wing and lateral wall are distinguished from the attachment muscle system on the lamellae and the existence of the calcareous covering; the outer surface of the wing is covered with a calcareous laver and the muscle attaches inside: the inner surface of the lateral wall is overlain



Text-fig. 10. The articulation of the upper and lower jaws of modern *Nautilus*. rh: rhyncholite, co: conchorhynch. Photos by Kanie. 1. Soft X-ray photograph of the jaw apparatus of *Nautilus pompilius* Linnaeus (natural size). Specimen JECOLN. P2  $\bigcirc$ . Loc. -150 m, cff Bindoy, Negros Oriental, Philippines. Condition: 30 seconds for an integral time, 40 kV-6 mA for an energy level per channel with Softex-EMB. 2. Lateral view of an articulated jaw of *Nautilus macromphalus* Sowerby (natural size). Specimen JECOLN. M5  $\bigcirc$ . Loc. off Nouméa, New Caledonia.

by a calcareous layer and underlies the muscle attachment.

It is suggested that the articulation and muscle system of both upper and lower jaws of tetragonitids are quite different from those of living cephalopods. The attachment areas of the muscles of tetragonitids are inferred from the morphological features of the muscle-scar, though there is no information on the jaw musclescar. The muscle attachments of the lower jaw in tetragonitids might be inside of the wing (Text-fig. 12), because of the existence of a calcareous layer on the wing surface. The muscle of the upper jaw also attaches inside of the wings

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Text-fig. 11. Anterior view of articulated upper and lower jaws of modern *Nautilus* species with jaw muscles (horizontal lines). U: upper jaw, L: lower jaw.



Text-fig. 12. Anterior view of articulated upper and lower jaws of a tetragonitid ammonite. Horizontal lines show the presumed jaw muscular system. U: upper jaw, L: lower jaw. Discussed with Y. Fukuda and K. Tanabe.

based on the similar lamellar structure. Therefore, the buccal mass system and morphological feature of the upper and lower jaws suggest an articulation system of the tetragonitid ammonite jaws as shown in Text-figure 13.

The feeding habits of tetragonitid ammonites which have a conchorhynch at the rostral part of the lower jaw are partly inferred from those of *Nautilus* with similar conchorhynch. Dean (1901), Bidder (1962), Haven (1972), Mikami & Okutani (1977), Kanie & Tanabe (1979),



Text-fig. 13. Restored jaw apparatus of tetragonitid ammonite within a buccal mass. The apparatus is covered with buccal membrance and papillae (BMP). Discussed with Y. Fukuda and K. Tanabe.

and others observed the feeding habits of living Nautilus. In their views, Nautilus mostly scavenges animals on the sea floor and rarely feed on slow moving benthos. The function of the rhyncholites and conchorhynchs to bite and cut up a prey has already demonstrated by Tanabe *et al.* (1980) through the observations of crop content remains. Nautilus is a slow moving animal, its maximum swimming speed being 25 cm/sec (Ward *et al.*, 1977), and usually 2.8 cm/sec (Mikami *et al.*, 1980). The tetragonitid animals with similar shells to Nautilus are probably also slow swimmers.

#### Conclusions

The general morphology of the jaw apparatuses of Late Cretaceous *Tetragonites* and *Gaudryceras* is described. They are of anaptychus-type, and are often preserved *in situ* within the body chambers. The upper and lower jaws are distinguished from each other by differences of morphology, size, internal structure and others. The tetragonitid ammonite jaws are more similar to those of modern *Nautilus* than to those of coleoids especially in the existence of calcareous coverings at the rostral part.

The jaw lamellae of tetragonitid ammonites are composed of carbonate apatite, and the rostra grade into wing surface layer consisting of calcite, whereas the lamellae of modern cephalopods consist of chitin. The rostra and denticles of *Nautilus* jaws are made up of aragonite and calcite, respectively. The outer lamellae of modern *Nautilus* are overlain by very thin aragonite layers. Therefore, the ammonite jaws originally seem to have been analogous to modern cephalopod jaws in regard to mineral composition and internal structure.

Based on the similarities of jaws and shell morphology, tetragonitid ammonites are interpreted to be analogous to *Nautilus* in living habits. The jaw apparatuses of tetragonitid ammonites are relatively large and probably stronger in comparison with those of *Nautilus*.

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#### Explanation of Plate 39

Figs. 1-6. Jaw apparatuses of Gaudryceras species.

1a, b. Upper jaw probably attributed to Gaudryceras sp. a: ventral, b: lateral views. ×1. MN6. Loc. the upper part of the Yubarigoenosawa, the Ikushumbetsu River, Ikushumbetsu area, central Hokkaido. Upper Yezo Group (Coniacian ?). M. Nihongi coll. (Kanie et al., 1978) 2. Lateral view of the upper jaw attributed to Gaudryceras sp. ×1. UH3193. Loc. the Kikumezawa, the Ikushmbetsu River, Ikushumbetsu area. Upper Yezo Group (Santonian ?). R. Saito coll. 3. Mode of occurrence of the upper jaw in Gaudryceras denseplicatum  $\times 1$ . GK. H8087. Loc. R4041p2, the upper part of Obirashibe River, Obira area, (Jimbo). nortwestern Hokkaido. Upper Yezo Group (Lower Santonian). H. Hirano and K. Tanabe coll. 4. Upper view of the deformed lower jaw attributed to Gaudryceras sp. ×2. GK. H8059. Loc. the upper part of the Kamikinenbetsu River, Obira area. Middle Yezo Group (Upper Turonian). T. Matsumoto coll. 5a-c. Upper jaw of Gaudryceras tenuiliratum Yabe. a: mode of occurrence, b: lateral, c: anterior views.  $a(\times 1)$ , b,  $c(\times 2)$ . UH3191. Loc. the Ikushumbetsu River, 200 m east from the junction with the Kikumezawa, Ikushumbetsu area. Upper Yezo Group (Santonian). R. Saito coll. (Nagao, 1931a, c; Tanabe et al., 1980). 6a, b. Lower jaw probably attributed to Gaudryceras denseplicatum (Jimbo). a: dorsal, b: lateral views, co: conchorhynch. ×1 YCM.Ur307002. Loc. U307p, Ikandai Valley, Urakawa area, southerncentral Hokkaido. Upper Yezo Group (Lower Campanian). Y. Kanie coll. (Kanie et al., 1978; Lehmann et al., 1980)

Photos, without whitening, by Kanie (1-3, 5, 6) and K. Tanabe (4).



5 a

5 c

**6** b

S. 49-55, Taf. 11.

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Ashibetsu 芦別, Furenai 振内, Hakkinzawa 白金沢, Ikandai 井寒台, Ikushumbetsu 幾春別, Juhachirinpan-Ichinosawa 18林班一の沢, Kamikinenbetsu 上紀念別, Kaneobetsu 金尾別, Kikumezawa 菊面沢, Naibuchi 内淵, Nakakinenbetsu 中紀念別, Obira 小平, Obirashibe 小平薬, Oyubari 大夕張, Santan 三丹, Shiyubari 主夕張, Souashibetsu 双芦別, Takishita 滝下, Tombetsu 頓別, Urakawa 浦河, Yubarigoenosawa 夕張越の沢

現生頭足類顎器との比較研究によるテトラゴニテス科アンモナイトの顎器:北海道・樺太 の白亜紀チューロニアン~マーストリヒシアン階産の石灰質ノジュール中に保存された Tetragonites glabrus, Gaudryceras denseplicatum, G. tenuiliratum, G. sp. の顎器の形 態と鉱物組成を記載した。これらの顎器は、外ラメラ・わい小化した内ラメラと石灰質の嘴状 部より構成されている。前者はカーボネイトアパタイト、後者はカルサイトよりなり、現生頭 足類の顎器と比較検討することによって、それぞれキチンとアラゴナイトより置換されたと考 えられる。石灰質嘴状部は、内部構造・鉱物、化学組成から、現生オウムガイ類のリンコライ ト(上顎)・コンコリンク(下顎)に酷似する。これらの資料に基づいて、テトラゴニテス科 アンモナイトの顎器と口球の一部を復元した。これらのアンモナイトの顎器は、オウムガイ類 よりがん丈で、住房部の大きさに対する顎器の大きさの比はより大きい。この事実は、テトラ ゴニテス科アンモナイトがオウムガイ類と類似の食生活をしていたことを示唆する。

蟹江康光

#### Explanation of Plate 40

Figs. 1-5. Jaw apparatuses of Tetragonites glabrus (Jimbo).

1a, b. Mode of occurrence (a) and lateral views (b) of upper jaw.  $a(\times 1), b(\times 2)$ . GK423. Loc. Sakhalin. Horizon unknown (Upper Cretaceous). collector unknown. 2. Lateral view of upper jaw; white calcareous covering is observable.  $\times 1.5$ . Loc. the lower part of the Kaneobetsu, Oyubari area, central Hokkaido. Middle Yezo Group (Turonian). Y. Kawashita coll. (Tanabe *et al.*, 1980). 3a, b. Mode of occurrence of lower and upper jaws (a), dorsal view of lower jaw (b). 1j: lower jaw, uj: rostral part of upper jaw.  $a(\times 3/4), b(\times 2)$ . GK.H2073. Loc. N143a, the lower part of the Santan, Naibuchi area, south Sakhalin. Upper part of the Miho Group (Lower Santonian). T. Matsumoto coll. (Kanie *et al.*, 1978; Tanabe *et al.*, 1980; Lehmann *et al.*, 1980). 4a, b. Lateral (a) and deformed dorsal (b) views of the lower jaw.  $\times 2$ . GK.H8049. Loc. R2251b, the middle part of the Nakakinenbetsu River, Obira area, northwestern Hokkaido. Saku Formation (Upper Turonian). H. Hirano and K. Tanabe coll. 5a-c. Mode of occurrence (a), lateral (b) and dorsal (c) views of the lower jaw.  $a(\times 1), b, c(\times 2)$ . GK. H8043. Loc. N382, the Juhachirinpan-Ichinosawa, the Naibuchi River, Naibuchi area. Upper part of the Miho Group (Santonian). T. Matsumoto coll. (Tanabe *et al.*, 1980).

Photos, without whitening, by K. Tanabe (1b, 3-5), I. Obata (1a) and Kanie (2).



4 b



5 c

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#### 741. EARLY JURASSIC PLANTS IN JAPAN. PART 4\*

#### TATSUAKI KIMURA and MASANORI TSUJII

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Abstract. This is the fourth of our serial papers on the early Jurassic plants in Japan. In this paper we describe *Ctenozamites sarrani* belonging to Pteridospermales, and *Anomozamites* sp. A, *Otozamites fujimotoi*, O. neiridaniensis sp. nov., *Pterophyllum* ex gr. *propinquum*, P. sp. A, *Ptilophyllum* cfr. *cutchense*, P. *nipponicum* sp. nov., P. *shinadaniense* sp. nov., P. sp. A. and P. sp. B, all belonging to Bennettitales.

The occurrence of *Ctenozamites* and *Ptilophyllum* is the first record not only in the early Jurassic flora but also in the older Mesozoic floras in Japan.

#### Introduction and acknowledgements

In continuation of our previous papers (Kimura and Tsujii, 1980a, b, '81), we here describe the fossil plants belonging to the Pteridospermales and Bennettitales from the Lower Jurassic Kuruma Group and Iwamuro Formation. The details of stratigraphy of plant-beaing formations and localities were given in our first paper (Kimura and Tsujii, 1980a).

We express our sincere gratitude to Professor Emeritus Thomas M. Harris, F. R. S. of the University of Reading for his helpful suggestions and [critical reading over the present manuscript. We also give our thanks to Mr. Gumpei Mori, Mr. Noriyuki Sasaki and Mr. Akihiro Kobayashi for their kind help in collecting the fossils here described.

#### Systematic description

Pteridospermales

Genus Ctenozamites Nathorst 1886: 122 Ctenozamites sarrani (Zeiller) Harris

Pl. 41, Fig. 1; Text-fig. 1

- Ctenopteris sarrani Zeiller: Zeiller, 1903, p. 53, pls. 6-7, fig. 1, la; pl. 8, figs. 1-2 (Hongay Coal-Field, North Viet-Nam); p. 292, pl. 44, figs. 3-4 (Tai-pin-tchang, North Yunnan; corresponding possibly to the Yipinglang Group): Sze, 1956, pp. 39, 146, pl. 35, figs. 3, 3a, 4 (Upper Triassic Yanchang Group, N-China).
- Ctenozamites sarrani (Zeiller) Harris: Harris, 1961, p. 170 (name): Sze et al., 1963, p. 198, pl. 58, fig. 1 (reinserted from Sze, 1956, pl. 35, fig. 3a); pl. 59, figs. 2, 3 (from Sze, 1956, pl. 35, figs. 3, 4): Li P.C., 1976, p. 126, pl. 40, figs. 5, 6 (Upper Triassic Yipinglang Group).

*Material*: NNW-146, 846 (counterpart) (Nishi-Neiridani).

Description: Pl. 41, Fig. 1 shows two pinna fragments. Leaf-outline and forking of main rachis are unknown. The pinnae are more than 14 cm long and 5.2 cm wide, nearly parallel-sided and with a thick

<sup>\*</sup> Received June 13, 1981; Read Jan. 23, 1981 at Sendai.

pinna rachis, 4 mm wide with longitudinal ribs on its surface. Segments (or pinnules) are typically rhomboidal, 2.9 cm long and 1.7 cm wide, and arising on the upper side of pinna rachis at a wide angle. The apex of segments is obtuse or rounded, its margins are entire and are flat but apical margins are sometimes reflexed. The bases of segments are not contracted. Veins are 10-11 in number at base, then forking once or twice, but the marginal ones are mostly simple. They are nearly parallel and end in the outer margin. Text-fig. 1 shows the form of typical pinnules and venation. The cuticle is not preserved.

Distribution and occurrence: In the Kuruma Group, a single specimen was obtained from the Negoya Formation. This species has been known from the Upper Triassic of North Viet-Nam and of North and Middle China.

*Remarks*: Our leaf corresponds in size and outline of segments (pinnules) to *Ctenozamites sarrani*. But it is somewhat different from Zeiller's original specimens of *Ctenopteris sarrani*. Differnces are: In the original specimens, the veins are more crowded and generally forking once, and the segments arise at reduced angle and are usually directed forwards, instead of arising at wide angle in ours.

Our leaf resembles a well-known European species, *Ctenozamites cycadea* (Berger) Schenk (references, see Harris, 1961, p. 169; 1964, p. 95). But our leaf differs from *Ctenozamites cycadea* in its large-sized segments with denser veins forking twice, instead of medium-sized segments (2 cm long and 1 cm wide, according to Harris, 1961, '64) with coarser veins (usually 6) forking once-twice in *C. cy-cadea*.

*Ctenozamites cycadea* finely illustrated by Delle (1967) from the Middle Jurassic of Tkvarchelian Coal-Basin (Transcaucasia) and *Ctenopteris cycadea* (Brongniart) by Yang (1978) from the Upper Triassic Daqiaode Formation, Sichuan, are also similar in form of segments to ours, but they are distinguished from ours by their medium-sized segments with coarser veins forking once or twice. *Ctenopteris megaphylla* described by Yang (1978) from the Upper Triassic Daqiaode Formation resembles *Ctenozamites sarrani* in form but smaller.

*Ctenozamites usnadzei* originally described by Doludenko and Svanidze (1969) from the Callovian of Georgia and later by Barnard and Miller (1976) from the Middle Jurassic of Iran, is distinguished from ours by its rhomboidal segments typically with a free acroscopic margin and decurrent basiscopic margin, and further by its small-sized segments with 1-3 small teeth at apex.

Recently 6 new *Ctenozamites* species were described from the Upper Triassic of China by Hsü et al. (1975) and Chen and Duan (1979a, b). Among them *Ctenozamites otoeis* described by Chen and Duan (1979a) from the Upper Triassic of Yanbian, has segments of similar size to our specimen of *C. sarrani*, but they are longer and have a decurrent base. The veins are more crowded.

This is the first record of the genus *Ctenozamites* in Japan.

#### Bennettitales

- Genus Anomozamites Schimper, 1870: 140 Anomozamites sp. A
- Anomozamites sp.: Kimura, 1959a, p. 22, pl. 8, fig. 3 (Iwamuro).
  - Material: A-0558 (a single specimen).

*Remarks*: Our leaf-fragment resembles Anomozamites major (Brongniart) Nathorst, and also *Pterophyllum pinnatifidum* Harris. Further material is needed for specific determination.

### Genus Otozamites Braun, 1842 Otozamites fujimotoi Kimura

#### Text-figs. 2a-b

*Otozamites fujimotoi* Kimura: Kimura, 1959a, p. 23, pl. 12, figs. 2-4, 6; text-figs. 12-13 (Iwamuro).

Material: Lectotype; A-4090 (Iwamuro). Paralectotypes; A-2037, 2054, 2081. Other specimens; A-0432, 0489, 0491, 2053, 2064-2066, 2068, 2088, 2091, 4091, 4094. Stratum typicum: Iwamuro Formation. Locus typicus: Iwamuro, Shirasawa-mura, Tone-gun, Gumma Prefecture. Derivatio nominis: After Professor Emeritus H. Fujimoto of the Tokyo University of Education.

Emended diagnosis: (Length and apex of leaf unknown.) In middle region rachis 2 mm wide, lamina 6 cm wide but narrower in basal part of leaf. Pinnae varied in form according to their position on a leaf, attached perpendicularly to the upper side of rachis by the lower half of base, thus the upper half free, forming a distinct angle entirely covering the rachis and the lower part of opposite pinna base. Pinnae mostly elongated but the proximal ones broadly deltoid or circular in form. Elongated pinnae 3 cm long or more, nearly parallel-sided, typically 1 cm wide. Deltoid or circular pinnae 8-20 mm long or more and 10-14 mm wide, with pointed apex. Circular pinnae with rounded apex restricted to basal part of a leaf. Veins on the lower half of pinna mostly simple or once forked and parallel, but those on free angle radiating, simple or once forked. Concentration of veins about 20 per cm. (Cuticle not preserved.)

Distribution and occurrence: Otozamites fujimotoi is occasional in the Iwamuro Formation, and seems to be local.

Comparison and remarks: Since Kimura

(1959a) described this species, only a few additional specimens were obtained. *Oto-zamites fujimotoi* is characterized by its varied pinnae in form, with distinct acroscopic basal angle.

Otozamites gissaricus illustrated by Iminov (1976) from the Middle Jurassic of Central Asia, is similar in form to ours, but O. gissaricus is a smaller leaf with narrower pinnae, especially on the proximal part of a leaf.

Specimens determined by Iminov (1976) from the Upper Triassic of Central Asia as *Otozamites pamiricus* Prynada are distinguished from *O. fujimotoi* by its pinnae with rounded or truncated apex. *Otozamites ptilophylloides* originally described by Barnard and Miller (1976) from the Middle Jurassic of Iran is similar in pinna form to *O. fujimotoi*, but it is macroscopically distinguished from ours by its pinnae mostly with rounded apex and smaller number of veins in each pinna.

Although we still do not know the form of the whole leaf of *Otozamites fujimotoi*, we feel sure it is distinct for its varied and characteristic pinna forms with distinct acroscopic angle.

Otozamites neiridaniensis Kimura and Tsujii sp. nov.

Pl. 42, Figs. 1-5; Pl. 43, Figs. 1-2; Text-figs. 3a-f

Otozamites molinianus Zigno: Oishi, 1940, p. 333, pl. 31, figs. 3, 3a, 4 (Neiridani).

Otozamites sp.: Oishi, 1940, p. 336, pl. 30, fig. 5 (Neiridani).

*Material*: Holotype; NEE-095 (Higashi-Neiridani). Paratypes; NEE-065, 174, TOB-013A, B, 050B, 050B', 050C (Tobiiwadani), NNW-084 (Nishi-Neiridani), NEG-251 (Negoya). Other specimens; NE-026 and other 9 specimens (Neiridani), NEE-007 and other 64, NNW-002 and other 57, NEG-006 and other 49, DAI-004 and other 38 (Dairagawa), TOB-



004 and other 21.

Stratum typicum: Negoya Formation, Kuruma Group. Locus typicus: Higashi-Neiridani, Asahi-machi, Shimo-Niikawa-gun, Toyama Prefecture (see Kimura and Tsujii, 1980a, p. 342, fig. 2). Derivatio nominis: After the Neiridani where the holotype and many other specimens were collected.

Diagnosis: Leaf rather large, long and narrow, linear; length estimated at 30 cm or more, width over the middle leaf typically about 2-2.5 cm, occasionally 3.7 cm. Leaf tapering gradually to both ends. Rachis rather thin, 1 mm wide, usually concealed. Pinnae oval or short-rectangular in form, typically 1.2 cm long and 0.8 cm wide, with rounded or obtusely pointed apex, attached on the upper side of rachis by the lower two-thirds or three-fourths of their width (accordingly upper onethird or one-fourth of acroscopic basal part free), at an angle of 60-90 degrees. Acroscopic basal corner slightly auriculated but seldom covering the pinna on the opposite side of the rachis. Basiscopic basal angle slightly contracted or occasionally straight but not decurrent. Near the leaf base, pinnae becoming shorter and more deltoid but still broad, and near the leaf apex, pinnae becoming narrower. Veins arising from proximal two-thirds or three-fourths of pinna base and radiating into the acroscopic corner and the rest of pinna, branching at all levels, at a concentration of 15 per cm at base. (Petiole not known and cuticle not preserved.)

Distribution and occurrence: Otozamites neiridaniensis is abundant in the Negoya Formation and rather rare in the Shinadani Formation. A doubtful leaf-fragment is known from Neiridani.

Comparison and discussion: Otozamites neiridaniensis is characterized by its oval or short-rectangular pinnae with a rounded or obtusely pointed apex. Its auricle is developed slightly if at all and its veins are crowded.

Oishi (1940) described several leaf-fragments from Neiridani (exact locality not clear) as Otozamites molinianus. But our observation of his figures shows that they do not belong to Zigno's species Otozamites neiridaniensis, but to because in his pinnae auricles are scarcely In Otozamites molinianus, developed. auricles are distinctly developed (Zigno, 1883; Möller, 1903). Oishi (1940) also described a single leaf-fragment as Otozamites sp. from Neiridani. This specimen agrees possibly with the basal pinnae of Otozamites neiridaniensis.

Among 160 *Otozamites* species hitherto known from the Mesozoic plant-beds, the following species resemble ours in rectangular or triangular pinnae with rounded or obtusely pointed apex, and with not or only slightly expanded auricle:

Text-figs. 1-3. (All natural size)

- 1. Ctenozamites sarrani (Zeiller) Harris; Outline of pinnules and venation (NNW-146; counterpart of NNW-846 shown in Pl. 41, Fig. 1).
- 2. Otozamites fujimotoi Kimura. 2a: A leaf-fragment with elongated and auriculated pinnae (A-4091; counterpart of A-4090 shown in Kimura, 1959a, text-fig. 12e). 2b-b''': Deltoid and auriculate pinnae at the basal part of a leaf (A-2054; redrawn from Kimura, 1959a, text-fig. 12b, a, d, c respectively).
- 3. Otozamites neiridaniensis Kimura and Tsujii sp. nov. 3a: A medium sized leaf (NEE-095, holotype; Pl. 42, Fig. 1). 3b: Ditto, with apex (TOB-050B, paratype). 3c: A large-sized leaf (TOB-050B', paratype; Pl. 42, Fig. 2). 3d: Apical part of a broken leaf (TOB-013B, paratype). 3e-e': Pinnae, showing the venation (NEE-255). 3f: A large-sized pinnae, showing the venation (NNW-084, paratype; Pl. 43, Fig. 1).

Otozamites falsus Harris: Harris, 1949, '69; Middle Jurassic of Yorkshire. O. indosinensis Zeiller: Zeiller, 1903; Upper Triassic of North Viet-Nam. O. kondoi Oishi: Oishi, 1940; Upper part of Middle Jurassic Moné Formation, NE-Japan. O. parallelus Phillips: Harris, 1969; Middle Jurassic of Yorkshire. O. reglei (Brongniart) Saporta: Brongniart, 1828; Bartholin, 1894; Möller, 1903 (cfr. reglei); Lower Jurassic of Bornholm. O. simpsoni Harris: Harris, 1949 (regarded as O. feistmanteli Zigno), '69; Middle Jurassic of Yorkshire.



Text-fig. 4. *Pterophyllum* ex gr. *propinquum* Goeppert; varied leaf-forms and venation. 4a: (NNW-459). 4b: (NNW-666). 4c: (NNW-670). 4d: (NNW-850). 4e: (NNW-531). 4f: (NNW-743). 4g: (NNW-810). 4h: (NNW-539). 4i: (NNW-743). 4j: (NNW-356; Pl. 43, fig. 5). 4k: (NNW-657). 4l: (NNW-326). 4m: (NNW-314). 4n: (NNW-009). (All natural size)

These differ as follows: Otozamites falsus has mostly rhomboid pinnae with an acute apex. O. indosinensis, O. parallelus and O. simpsoni have more elongate pinnae. Ptilophyllum khargaense (Kilpper, 1966) from the Middle Jurassic (?) of Egypt has pinnae of rather similar shape but the lower ones differ in being imbricate. The closest to our species is Otozamites kondoi but its pinnae have a nearly symmetrical base and its veins which are coarser, only fork near their origin.

#### Genus Pterophyllum Brongniart, 1828: 25

*Pterophyllum jaegeri* Brongniart described by Oishi (1931) from Tsuchizawa is not *Pterophyllum* but referable to *Nilssonia* sp. A to be described later.

#### Pterophyllum ex gr. propinquum Goeppert

Pl. 41, fig. 2; Pl. 42, figs. 6-7; Pl. 43, figs. 3-6; Text-figs. 4a-n

Specimens from East Asia:

- Pterophyllum propinguum Goeppert: Oishi, 1931, p. 245, pl. 18, figs. 1-2 (Tsuchizawa, Kuruma Group): Yabe and Oishi, 1933, p. 227, pl. 34, fig. 2 (Lower-Middle Jurassic of Liaoning, NE-China): Oishi, 1940, p. 344 (remarks): Kimura, 1959b, p. 70, pl. 2, fig. 2 (Kotaki Coal-Field, Kuruma Group): Sze et al., 1963, p. 157, pl. 61, fig. 8 (reinserted from Yabe and Oishi, 1933, pl. 34, fig. 2).
- Pterophyllum aff. propinguum Goeppert: Yabe and Oishi, 1929, p. 91, pl. 18, fig. 6 (Daedong Group, N-Korea).
- For Europeaw references, see Antevs, 1919, p. 28.

*Material*: NNW-009 and other 32 specimens (Nishi-Neiridani), DAI-027 (Dairagawa), SHI-035, 040 (Shinadani), Kr-087, 090 (Tsuchizawa), 4962113 (Kotaki Coal-Field), A-7535, 751124 (Iwamuro).

Description: Leaves are varied in size and form, but are mostly medium-sized

and oblanceolate in outline. Leaves taper gradually towards the proximal end, are more than 6.5 cm long and 4 cm wide at the widest portion, and possibly contract suddenly at the apex. The rachis is rather thin, 2 mm wide. The petiole is not known. The pinnae arise at a wide angle from the lateral margins of rachis. They are rectangular, up to 2 cm long and are shorter towards the base of the leaf. The width of pinnae is variable but often about 1.2 cm; pinnae near the leaf apex are narrow. Margins of adjacent pinnae are mostly apart but are always connected to one another near the rachis by a web of lamina. The apex of pinnae is truncate in lower part of leaf, but rounded in the middle of leaf. Veins are parallel and simple or often fork once. Their concentration is typically 25 per cm at the middle of The cuticle is not preserved. pinna. Text-figs. 4a-n show varied pinna forms and venation.

Distribution and occurrence: This species is locally very abundant in the Negoya Formation, but rather rare in the Shinadani and Iwamuro Formations. In East Asia, this species occurs in the Lower-Middle Jurassic of NE-China and the Upper Triassic (?) of N-Korea.

*Remarks*: *Pterophyllum* ex gr. *prop-inquum* is characterized by its rectangular pinnae with an expanded base and truncated or rounded apex. Our leaves are indistinguishable from *Pterophyllum pro-pinquum* illustrated in detail by Antevs (1919) from the Liassic Hör Sandstone, Sweden, although in our large collection, leaves with elongated pinnae such as Antevs' pl. 3, figs. 10-12 have not been found.

The leaves regarded by Oishi (1931) as *Pterophyllum propinquum* from Tsuchizawa and by Yabe and Oishi (1929) as *P.* aff. *propinquum* from N-Korea are



indistinguishable from ours, although the Korean leaf is about twice as large as ours. A leaf-fragment regarded by Yabe and Oishi (1933) as *Pterophyllum propinguum* possibly from the Beipiao Group (Lower-Middle Jurassic), NE-China is similar in pinna form to ours.

*Pterophyllum* leaves with similar pinna form to ours have been described as follows:

Pterophyllum pinnatifidum Harris: Harris, 1932; Rhaetic of East Greenland. P. kochi Harris: Harris, 1926, '32; Ditto: Lundblad, 1950; Rhaeto-Liassic of Sweden. P. abnorme Etheridge: Walkom, 1917: Jones and Jersey, 1947; Upper Triassic of Queensland. P. sensinovianum Heer: Heer, 1876; Jurasso-Cretaceous of the upper Amur: Vakhrameev and Doludenko, 1961; Upper Jurassic of Bureja Basin.

They are distinguished from ours as mentioned below. In *Pterophyllum pinnatifidum*, pinna apices are often acuminate. In *Pterophyllum kochi*, pinna bases are not contiguous except in the apical part of leaf and are square in the middle of leaf, and slightly contracted in the proximal part of leaf. In ours, pinna bases are usually contiguous from one pinna to the next.

A little known *Pterophyllum abnorme* is distinguished from ours by its usually twice forked veins instead of simple or once forked ones in ours. In *Pterophyllum* sensinovianum, its pinnae are more elongated than those of ours.

Pterophyllum propinquum illustrated by Iminov (1976) from the Upper Triassic of Pamir is, in our opinion, rather close to *P. tsetzei* Schenk or to *P. subaequale* Hartz because of its large-sized leaves with more elongated pinnae.

Pterophyllum sp. A illustrated by Weber (1968) from the Rhaeto-Liassic of Bayreuth resembles *P. propinquum. Pterophyllum exhibens* originally described by Li, P. C. (1964) from the Upper Triassic Xujiahe Formation, Sichuan and later by Zhou, T. S. (1978) from the Upper Triassic Wenbinshan Formation, Fujian, China, is similar in its expanded pinna base, but it is distinguished by its more elongated pinnae with smaller number of veins (5-12) in each pinna.

#### Pterophyllum sp. A

Pterophyllum sp. A and sp. B: Kimura, 1959a, p. 24, pl. 9, figs. 2-3 (Iwamuro).

*Remarks*: Kimura (1959a) illustrated two *Pterophyllum* leaf-fragments as *P*. sp. A and sp. B. Our study convinced us that they are conspecific but the specimens are too poorly preserved to determine. Their cuticle is not preserved.

Text-figs. 5-6 (a-f). (All natural size, unless otherwise indicated)

Ptilophyllum cfr. cutchense Morris. 5a: A leaf-fragment, showing the venation (TE-001; Pl. 43, Fig. 7). 5b: Ditto, with apex (TE-003; Pl. 43, Fig. 8). 5b': Venation drawn from 5b.

<sup>6.</sup> Ptilophyllum nipponicum Kimura and Tsujii sp. nov. 6a: A typical leaf (NEG-720, holotype; Pl. 43, Fig. 9). 6a': Venation drawn from 6a. 6b: Lower part of a leaf (NNW-806, Pl. 41, Fig. 3). 6b': Possibly representing the basal part of 6b. 6c: Basal part of a leaf, bearing broader pinnae with truncated distal margin (NE-067, paratype). 6d: Apical part of a medium-sized leaf bearing rather remotely set pinnae (NNW-353, paratype; Pl. 41, Fig. 4). 6d': Venation drawn from 6d. 6e: A leaf-fragment bearing rectangular pinnae (NEE-074; Pl. 42, Fig. 8). 6e': Venation drawn from 6e. 6f: A leaf-fragment bearing broader and remotely set pinnae (NNW-745).

Genus Ptilophyllum Morris, 1840: 327

About 60 Mesozoic species have been described. Many of them have similar pinnae. Thus except the leaves with distinct morphological features, it is getting difficult to identify them depending only on their macroscopical forms. Our knowledge of the epidermis of this genus owes much to the studies of Jacob and Jacob (1954), Kilpper (1965, '66), Harris (1941, '42, '46, '49, '53, '69), Doludenko (1963), Bose and Kasat (1972) and Barnard and Miller (1976).

We recognized five forms in the Lower Jurassic of Japan. This is the first record



of this genus in the older Mesozoic floras in Japan. The phytogeographical implication on the occurrence of this genus in

the Japanese Lower Jurassic will be mentioned in our later paper treating the concluding chapter of this series.

#### Key to the Japanese Lower Jurassic species of Ptilophyllum

1.	Leaf large-sized, widest part over 5.4 cmP. nipponicum sp. nov
1′.	Leaf medium-sized about 3 cm wide
1″.	Leaf small-sized width less than 1 cm
2.	Pinnae elongate-rectangularP. shinadaniense sp. nov.
2′.	Pinnae elongate-oval or rectangularP. sp. A
3.	Pinnae elongate-triangularP. cfr. cutchense
3′.	Pinnae rhombicP. sp. E

Ptilophyllum cfr. cutchense Morris

Pl. 43, Figs. 7-8; Text-figs. 5a-b

Comparable specimens:

Ptilophyllum cutchense Morris: Bose and Kasat, 1972, p. 118, pl. 1, figs. 7-10; pl. 2, figs. 14-17 (Upper Jurassic of Rajmahal Hills and Lower Cretaceous Jabalpur Series, India).

*Material*: TE-001, 003 (Teradani, Negoya Formation).

Description: Two small-sized leaffragments were obtained. The leaves are 0.8-1 cm wide with an acuminate apex as shown in Pl. 43, Fig. 7. The rachis which is partially exposed, is about 0.5 mm wide. Pinnae are attached on upper surface of rachis, closely set or imbricate, mostly attached by entire base, arising at an angle of about 60-90 degrees. Pinnae are as a whole linear, elongate-triangular, sometimes rhomboidal, sublate or ovate in form; margins are straight. Apices are usually obtuse or rounded and the basiscopic basal margin straight, slightly contracted, rounded or decurrent. Veins arise from the whole base. They are parallel, mostly simple, 5 in number on each pinna. Cuticle is not preserved.

Distribution and occurrence: This form is rare and known only from Teradani, Negoya Formation.

*Remarks*: Our leaves reaves resemble closely some smaller leaves illustrated by Bose and Kasat (1972) as *Ptilophyllum cutchense*. *Ptilophyllum horridum* Roy and *P. institacallum* Bose are allied forms to *P. cutchense*. But both are distinguished from our leaves by their radiating

Text-figs. 6 (g-k)-9. (All natural size, unless otherwise indicated)

- 6. Ptilophyllum nipponicum Kimura and Tsujii sp. nov. 6g: A leaf-fragment bearing elongated pinnae with decurrent base and contracted acroscopic base (NNW-241, paratype). 6g': Vanation drawn from 6g. 6h: Small-sized leaf-fragment (NNW-353, paratype). 6i: Large-sized leaf-fragment bearing broader pinnae with slightly decurrent base (NNW-241, paratype). 6j: Apical part of a leaf (NE-169). 6k: Basal part of a leaf with thick petiole of which base is expanded. Pinnae are small-sized, and with straight or slightly contracted base (NNW-111, paratype; Pl. 43, Fig. 10).
- Ptilophyllum shinadaniense Kimura and Tsujii sp. nov. 7a: A part of leaf (SHI-007, holotype; Pl. 41, Fig. 5). 7b: A leaf-fragement, showing the venation (SHI-024, paratype). 7c: Ditto (SHI-019, paratype; Pl. 41, Fig. 6).
- 8. Ptilophyllum sp. A; a part of leaf (SHI-030; Pl. 43, Fig. 11). 8': Venation drawn from 8.
- Ptilophyllum sp. B. 9a: A part of leaf (NEG-51; Pl. 43, Fig. 12). 9b: Ditto (A-0379).
  9b': Venation drawn from 9b.

and forked veins instead of mostly simple and parallel ones.

Ptilophyllum cantherifera originally described by Douglas (1969) from the Lower Cretaceous of Victoria (Australia) is similar in form, but is distinguished by its pinnae attached by central basal area instead of by the entire base in ours.

Thus our leaves are most like the smallsized *Ptilophyllum cutchense*, but we here refrain from fully identifying them with this Indian species, because our leaves without cuticle are only two and incomplete.

#### Ptilophyllum nipponicum Kimura and Tsujii sp. nov.

Pl. 41, Figs. 3-4; Pl. 42, Fig. 8; Pl. 43; Figs. 9-10; Text-figs. 6a-k

*Material*: Holotype; NEG-270 (Negoya). Paratypes; NNW-111, 241, 353, 697, 806 (Nishi-Neiridani), NE-067 (Neiridani). Other specimens; NEG-071 and other 3 specimens, NNW-004 and other 6, NEE-074 (Higashi-Neiridani), NE-169, DAI-086 (Dairagawa), IN-001, 002 (Inumatadani).

Stratum typicum: Negoya Formation, Kuruma Group. Locus typicus: The Negoya, Asahimachi, Shimo-Niikawa-gun, Toyama Prefecture. Derivatio nominis: After Nippon meaning Japan in Japanese.

*Diagnosis*: Leaf large, oblong or oblanceolate in form; the widest part 5.4 cm. (Whole length and leaf apex unknown.) Rachis slender 2 mm wide at the middle of leaf but 3 mm wide at the basal part. Petiole about 5 cm long with striated surface, base distinctly expanded. Pinnae varied in form according to their position on a leaf, attached on the upper side of rachis at an angle of 50-90 degrees by entire base, set closely but sometimes remotely. Middle pinnae elongate-rectangular, typically 3.5-4 cm long and 0.5-1 cm wide. Apical pinnae similar in form to those of middle ones but smaller in size. Basal pinnae typically deltoid in form, 1.5 cm long and 0.5-0.8 cm wide at base, but smaller towards the base of leaf. Acroscopic basal margin mostly rounded and basiscopic basal margin straight, slightly contracted or slightly decurrent. Apex acutely or obtusely pointed. Veins numerous, divergent, forking at all levels, 28-30 per cm in density at the middle of pinna. (Cuticle not preserved.)

Distribution and occurrence: Ptilophyllum nipponicum is common only in the Negoya Formation.

*Comparison and discussion*: It is clear that our leaves belong to *Ptilophyllum* redefined by Harris (1969), because our acroscopic pinna base is not auriculated and the basiscopic pinna base is typically decurrent.

*Ptilophyllum nipponicum* is characterized by its large-sized leaf with elongated pinnae above and deltoid pinnae below with divergent and forked veins, and its thick petiole and expanded base.

Several *Ptilophyllum* species with large leaves have been described. *Ptilophyllum nipponicum* closely resembles *P. hirsutum* Thomas and Bancroft described in detail by Harris (1949, '69) from the Middle Jurassic of Yorkshire. But the pinnae of *Ptilophyllum hirsutum* do not have a decurrent basal margin.

Ptilophyllum harrisianum described by Kilpper (1968) from the Liassic of N-Iran has pinnae of similar size but with an auriculate acroscopic base. Kilpper's Ptilophyllum zirabense (1968) has more elongated pinnae.

Ptilophyllum shinadaniense Kimura and Tsujii sp. nov.

Pl. 41, Figs. 5-6; Text-figs. 7a-c

Material: Holotype; SHI-007 (Shinadani).
Paratypes; SHI-019, 024.

Stratum typicum: Shinadani Formation, Kuruma Group (see Kimura and Tsujii, 1980a, p. 342, fig. 2). Locus typicus: The Shinadani (near Agero Village), a branch of the Dairagawa, Omi-machi, Nishi-Kubiki-gun, Niigata Prefecture. Derivatio nominis: After the locality, the Shinadani where the holotype and paratypes were collected.

*Diagnosis*: Leaf medium-sized, 3 cm wide. (Whole length unknown.) Pinnae arising at an angle of 65-80 degrees to the rachis, elongate-rectangular in form, typically 1.8 cm long and 3.5-4 mm wide. Width of pinnae about one-fifth its length. Pinnae straight or very slightly falcate, apex obtusely pointed, basiscopic basal angle normally decurrent but often concealed by the pinna below. Veins arising from whole base of pinna, nearly parallel, often forking once at the middle of pinna, 8-10 in number in each pinna. (Cuticle not preserved.)

Distribution and occurrence: Ptilophyllum shinadaniense is rare and known only from the Shinadani Formation.

Discussion and comparison: Ptilophyllum shinadaniense is characterized by its medium-sized elongate-rectangular pinnae with parallel veins often forking once.

Ptilophyllum contiguum originally described by Sze (1949) and later by Wu et al. (1980) from the Lower Jurassic Xiangxi (formerly Hsiangchi) Formation in W-Hubei, China has pinnae of similar form. But it is distinguished by its smallsized leaf (13.5 cm long and 1.6 cm wide) and fewer veins (4-6 in each pinna).

Ptilophyllum basekgabense originally described by Barnard and Miller (1976) from the Middle Jurassic of N-Iran has smaller pinnae (8-15 mm long and up to 3 mm wide) with fewer veins (4-5 in each pinna).

Some leaves of *Ptilophyllum caucasicum* Doludenko and Svanidze from the Upper Jurassic of Georgia resemble *P. shinadaniense*, but others have very long and narrow pinnae (Doludenko and Svanidze, 1969).

Ptilophyllum boreale (Heer) (Seward, 1917; Zamites borealis by Heer, 1874) from the Lower Cretaceous of Greenland is distinguished by its stout rachis and fewer veins (3-5 in number in each pinna).

*Ptilophyllum sokalense* described by Doludenko (1963) from the Jurassic of W-Ukraine has pinnae of similar form, but their basiscopic margin is mostly straight or slightly contracted.

# Ptilophyllum sp. A

Pl. 43, Fig. 11; Text-fig. 8, 8' Material: SHI-030 (Shinadani).

Description: Leaf medium-sized, elongated, 2.5 cm wide over the most part; neither end known. Pinnae elongate-oval or rectangular, attached by entire base on the upper side of rachis (2 mm wide) at an angle of 60-90 degrees, 1.4 cm long and 0.5 cm wide. Acroscopic basal angles mostly rounded or slightly decurrent, apices rounded or obtusely pointed. Veins indistinct, possibly simple and parallel, not radiating, 9 in each pinna. (Cuticle not preserved.)

*Distribution and occurrence*: A single specimen was obtained from the Shinadani Formation.

*Remarks*: Our specimen is characterized by its elongate-oval or rectangular pinnae. Our specimen is distinguished from *Ptilophyllum shinadaniense* from the same locality by its shorter pinnae.

Some leaves regarded as *Ptilophllum caucasicum* Doludenko and Svanidze (Doludenko and Svanidze, 1969; Doludenko and Orlovskaja, 1976) from the Upper Jurassic of Georgia and Kazakhstan are similar in pinna form and venation to ours, but others have very long and narrow pinnae.

## Ptilophyllum sp. B

Pl. 41, Fig. 7; Pl. 43, Fig. 12; Text-figs. 9a-b'

Material: A-0379, 0456, 7564 (Iwamuro), NEG-051 (Negoya).

Description: Leaves small-sized, long and narrow and nearly parallel-sided, 4 mm wide. Rachis 1 mm wide but mostly concealed. Pinnae rhombic, attached by entire base at a wide angle on the upper side of rachis, closely set or imbricate, apices usually obtuse, acroscopic basal margin straight or rounded, basiscopic basal margin mostly straight but sometimes slightly contracted. Veins indistinct, arising from the whole base, parallel, probably simple, about 5 in each pinna. (Cuticle not preserved.)

Distribution and occurrence: This form occurs sparsely in the Iwamuro and Negoya Formations.

*Remarks*: Judging from the mode of pinna attachment and venation, our leaves may belong to *Ptilophyllum*. So far as we know, such small-sized leaves as ours have been described as *Ptilophyllum* or *Otozamites* as follows: *Ptilophyllum bengalense* (Oldham) Schimper: Kilpper, 1968; Liassic of N-Iran. *P. elongatum*  Douglas: Douglas, 1969; Lower Cretaceous of Victoria (Australia), P. kazachstanicum Genkina: Genkina, 1963; Middle Jurassic P. marginatum Proof Eastern Ural. svirjakova: Prosvirjakova, 1966; Middle Jurassic of Mangwishlak. P. minor Kilpper: Kilpper, 1968, Liassic of N-Iran. Otozamites abreviatus Feistmantel: Feistmantel, 1879 (pars): Jurassic of Madras Coast. O. bengalensis (Oldham and Morris) Schimper: Oldham and Morris, 1862 (regarded by them as Palaeozamia bengalensis): Jurassic of India. O. bunburyanus Zigno: Zigno, 1881; Middle Jurassic of Italy. O. contiguus Feistmantel: Feistmantel, 1876; Jurassic of India. O. kilpberiana Barnard and Miller: Barnard and Miller, 1976; Middle Jurassic of N-Iran. O. tenuatus (Leckenby) Harris: 1969; Middle Jurassic of Yorkshire. O. vemavaramensis Bose and Jain: Bose and Jain, 1967; Upper Jurassic (?) of India.

Among these, our leaves are most like *Ptilophyllum bengalensis* in pinna form. But we refrain from determining it specifically because our leaves are all incomplete.

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Barnard, P.D.E. and Miller, J.C. (1976):

Explanation of Plate 41

Fig. 1. Ctenozamites sarrani (Zeiller) Harris: A part of leaf, with two pinna fragments. Judging from their wedge-shaped arrangement, they may represent the pinnae forking dichotomously below (NNW-846; counterpart of NNW-146 partly shown in Text-fig. 1), ×1.

Fig. 2. Pterophyllum ex gr. propinquum Goeppert; (NNW-695), ×1.

Fig. 7. Ptilophyllum sp. B; a leaf-fragment (A-7564), ×2.

Figs. 3-4. Ptilophyllum nipponicum Kimura and Tsujii sp. nov.; both, ×1. 3. Lower half of a leaf (NNW-806, paratype; Text-fig. 6b). 4. Apical part of a leaf, with rather remotely set pinnae (NNW-353; Text-fig. 6d, d').

Figs. 5-6. *Ptilophyllum shinadaniense* Kimura and Tsujii sp. nov.; both, ×1. 5. (SHI-007, holotype; Text-fig. 7a). 6. (SHI-019, paratype; Text-fig. 7c).



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## Explanation of Plate 42

- Figs. 6-7. *Pterophyllum* ex gr. *propinquum* Goeppert; both, ×1. Leaf-fragments. 6; (NNW-073), 7; (NNW-600).
- Fig. 8. *Ptilophyllum nipponicum* Kimura and Tsujii sp. nov.; apical part of a leaf (NEE-074; Text-fig. 6e), ×1.

Figs. 1-5. Otozamites neiridaniensis Kimura and Tsujii sp. nov.; all figures, ×1. 1, 3. Leaves of typical size. 1; (NNE-095, holotype; Text-fig. 3a), 3; (NE-046). 2. A large-sized leaf (TOB-050B', paratype; Text-fig. 3c). 4. Basal part of a leaf (NEE-250). 5. Ditto (NEE-073).



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Geographical names in Japan: Agero 上路, Dairagawa 大平川, Higashi-Neiridani 東寝入谷, Inumatadani 犬俣谷, Iwamuro 岩室, Kita-Otari 北小谷, Kotaki 小滝, Kuruma 来馬, Moné 舞 根, Nariwa 成羽, Negoya 似虎谷, Neiridani 寝入谷, Nishi-Neiridani 西寝入谷, Shinadani 榀谷, Teradani 寺谷, Tobiiwadani 飛岩谷, Tsuchizawa (Tunzawa) 土沢, Geographical names in China (with former expression in parentheses): Beipiao (Peipiao) 北票, Daqiaode 大蕎地, Fujian 福建, Hubei (Hupeh) 湖北, Liaoning 遼寧, Sichuan (Szechuan) 四川, Taipingchang (Tai-ping-tchang) 太平場, Wengbinshan 文賓山, Xujiahe (Hsuchiaho) 須家河, Xiangxi (Hsiangchi) 香渓, Yanbian 塩辺, Yanchang (Yenchang) 延長, Yipinglang 一平浪, Yongren (Yungjen) 永仁, Yunnan 雲南, Geographical name in Korea (with former expression): Daedong (Taedong) 大同

日本の初期ジュラ紀植物化石。その4:前編につづき,来馬層群および岩室累層からつぎ のシダ種子およびペネチテス目に属する植物化石を記載した。

シダ種子目: Ctenozamites sarrani (Zeiller) Harris

ベネチテス目: Anomozamites sp. A, Otozamites fujimotoi Kimura, O. neiridaniensis Kimura and Tsujii sp. nov., Pterophyllum ex gr. propinquum Goeppert, P. sp. A, Ptilophyllum cfr. cutchense Morris, P. nipponicum Kimura and Tsujii sp. nov., P. shinadaniense Kimura and Tsujii sp. nov., P. sp. A, P. sp. B.

以上のうち, Ctenozamites 属は、日本で最初の産出であり、Ptilophyllum 属は、日本の古期中生界からの最初の産出である。とくに後者の産出は、日本の初期ジュラ紀植物群の古植物地理的位置を判断する上で重要な資料となる。 木村達明・辻井正則

# Explanation of Plate 43

- Figs. 1-2. Otozamites neiridaniensis Kimura and Tsujii sp. nov.; both, ×1. 1. A large-sized leaf-fragment with elongated pinnae (NNW-084, paratype; partly shown in Text-fig. 3f). 2. Basal part of a leaf (TOB-032).
- Figs. 3-6. Pterophyllum ex gr. propinguum Goeppert; all figures, ×1. Leaf-fragments. 3; (NNW-240), 4; (NNW-356B), 5; (NNW-356; Text-fig. 4j), 6; (NNW-014).
- Figs. 7-8. Ptilophyllum cfr. cutchense Morris; both, ×1. Unsatisfactorily preserved leaf-fragments. 7; (TE-001; Text-fig. 5a), 8; (TE-003; Text-fig. 5b).
- Figs. 9-10. Ptilophyllum nipponicum Kimura and Tsujii sp. nov; both, ×1. 9. A leaf lacking both apical and basal portions (NEG-270, holotype; Text-fig. 6a, a'). 10. Basal part of a leaf with petiole of which base is expanded; upper arrow indicates a thick rachis and lower one an expanded petiole-base (NNW-111, paratype; Text-fig. 6k).
- Fig. 11. Ptilophyllum sp. A; a leaf-fragment (SHI-030; Text-fig. 8), ×1.
- Fig. 12. Ptilophyllum sp. B; a leaf-fragment (NEG-051; Text-fig. 9a), ×2.

KIMURA and TSUJII: Early Jurassic plants



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# 742. JAPANESE MIOCENE CIRRIPED BALANUS SENDAICUS: A COMPARISON WITH TETHYAN BALANUS CONCAVUS GROUP\*

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Abstract. Balanus sendaicus was described as a new species by Hatai, Masuda and Noda (1976) based on a single scutum from the Miocene Moniwa Formation, northern Honshu, Japan. However, the scutum deposited in the Saito Ho-on Kai Museum, Sendai, Japan (Hatai et al., 1976) could not be found. Tergum and shell wall, which with scutum comprise the skeletal elements, had not been described. According to my investigation of topotype materials and of materials illustrated by Hatai et al. (1976), the shell walls described and illustrated under the name of B. rostratus by Hatai et al. (1976) belong to B. sendaicus. B. sendaicus, including tergum undescribed by Hatai et al. (1976), is therefore defined in this paper. B. sendaicus is distinct from the already known taxa, as mentioned by Hatai et al. (1976). However, B. sendaicus is closely related to the extinct B. concavus sensu de Alessandri (1906), which flourished in the Mediterranean region (Tethys) from the Oligocene to Pleistocene. B. sendaicus belongs to the B. concavus species group represented by the nominal species based on having transverse septa in longitudinal (parietal) tubes of shell wall and the characteristic ornamentation of scutum. Disappearance of the B. concavus group from the Atlantic and the Mediterranean regions in Pleistocene is mysterious as mentioned by Pilsbry (1916), though "four" species belonging to the B. concavus group are living in the eastern Pacific (Newman, 1979). B. sendaicus is restricted to the Miocene, is the first fossil record of the B. concavus group from the east Asia, and indicates the presence of a connection between the Tethys and the east Asia.

#### Introduction

Balanus sendaicus was proposed as a new species by Hatai, Masuda and Noda (1976), based on a single scutum from the Miocene Moniwa Formation of the Sendai area of northeast Honshu, Japan. Seven new species including *B. sendaicus*, an already known species of *B. rostratus*, and an indetermined species of *Balanus* were described from the same locality. Four new species including *B. sendaicus* were based on scuta and three new species, *B. rostratus* and *B.* sp. based on the shell walls. No tergum was described.

Fossil specimens are usually found as fragmental skeletal elements of shell wall and opercular valves, and therefore it is difficult to reconstruct the exoskeleton of a species from the mixture of fragments of several extinct species. Under these circumstances, each fragmental skeletal element, such as shell wall, scutum or tergum, has been frequently chosen as

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representative of a new species. Thus in some cases, an individual consisting of three kinds of skeletal elements, may possesses three different species names. Therefore, biological study of fossil specimens of extinct species is faced with difficulties unless basic work, revealing the precise combination of skeletal elements, is done first. When a complete shell wall is discovered, it is necessary to determine whether or not opercular valves are preserved in the shell cavity.

In this paper the following points are considered: 1) Description of shell wall and tergum of *B. sendaicus*, whose original description was based on a single scutum, 2) The status of *B. sendaicus* as a previously unknown taxon, as discussed by Hatai *et al.* (1976), 3) The affinity of *B. sendaicus* with known species and species groups, and 4) The stratigraphic and geologic distributions in Japan and the paleoecological significance of *B. sendaicus*.

# Balanus sendaicus, type-locality and formation

The type-locality of this species is an outcrop (Loc. 4) of the Miocene Moniwa Formation situated about 200 m south of the Oide Bridge across the Natori River, near Moniwa, Sendai City, Miyagi Prefecture, northeast Honshu, Japan. The Miocene Moniwa Formation overlies the Miocene Takadate Andesite with a clino-unconformity. B. sendaicus was collected from the basal part of the Moniwa Formation which consists of gravelly calcareous medium-grained sandstone containing very abundant well preserved marine fossils. Reports on brachiopods, shark teeth, fossil problematica, and on pelecypods have already been published (Hatai et al., 1973, 1974a, 1974b, 1974c, respectively). Α report on cirripeds was the fifth of this series from the same locality. Besides B. sendaicus, shell walls but no terga of B. rostratus, B. miyagiensis, B. tohokuensis, B. rikuzenensis, and B. sp. and scuta of B. oidensis, B. moniwaensis, and B. natorianus were described. Except for B. rostratus and B. sp., all were proposed as new species.

# Combination of skeletal elements of B. sendaicus

The writer has collected many well preserved fossil cirripeds from the above locality. Four species by shell walls, five species by scuta, and two species by terga were distinguished. However, the combination of these exoskeleton elements was not clear. Judging from shell size and the frequency of occurrence of each element, it was presumed that shell wall of *B. rostratus* and scutum of *B. sendaicus*. as described by Hatai et al. (1976) and a previously undescribed tergum probably constituted skeletal elements of one species. This is also supported by the fact that the following morphological characteristics of shell walls of B. rostratus identified by Hatai et al. (1976) are clearly different from the true B. rostratus proposed by Pilsbry (1916); narrow radii and oblique summits of radii. In order to confirm the above conjecture additional material was sought.

Recently some shell walls collected from the same locality, identical with *B. rostratus* of Hatai *et al.* (1976), were found to contain paired scuta and terga within their shell cavity. The specimen of *B. rostratus* illustrated by Hatai *et al.* (1976, pl. 2, fig. 2, left, and Pl. 44, Figs. 2a-c, left, in this paper) was also found to contain four opercular valves in its shell cavity (Pl. 44, Figs. 2d-g, in this paper). It is important to note that morphologically these scuta are identical with the holotype scutum of *B. sendaicus* (Hatai *et al.* 1976, pl. 2, figs. 9 and 10). It can, therefore, be concluded that *B. sendaicus* is conspecific with *B. rostratus* of Hatai *et al.* (1976). However, *B. rostratus* of Hatai *et al.* (1976) is not the true *B. rostratus* in the sense of Pilsbry (1916) and the name *B. sendaicus* should be used. As a result of the confirmation of the correct combination of skeletal elements, the taxonomic status and the affinity of *B. sendaicus* can be discussed.

In the description of B. sendaicus, Hatai et al. (1976) pointed out that "In the external sculpture, the present scutum resembles that of Balanus concavus Bronn (Pilsbry, 1916, p. 100, pl. 21, fig. 1), a Pliocene species of the British Red Crag especially in cancellated sculpture, bifurcation of the longitudinal striae, general shape of scutum,...". Thus, owing to distinct growth ridges and longitudinal ridges or striations frequently bifurcated on external surface, scutum of B. sendaicus possesses common morphological characteristics found in scuta of all members of the B. concavus species group. The tergum of B. sendaicus described here also has a long spur and partially closed spur furrow, characteristics found in some members of the *B. concavus* group. These morphological characteristics found in B. sendaicus are also present in members of the *B. perforatus* species group. However, in members of the *B. concavus* group, the external surface of scutum is heavily ornamented, the aperture is small, and the carina strongly curved. Therefore, B. sendaicus is regarded as a member of the *B. concavus* group.

# Morphological comparisons of *B.* sendaicus with the living taxa of the *B.* concavus group

According to Newman and Ross (1976), the B. concavus group contains the following "five" living taxa, B. aquila (=B.gregarius (Conrad), 1856; a senior synonym of B. aquila Pilsbry, 1907, see Zullo, 1964:360, 1969:6, 1979:4), B. concavus concavus, B. c. mexicanus, B. eyerdami (=B. panamensis Rogers, 1948; a seniorsynonym of B. eyerdami Henry, 1960, D. P. Henry, pers. comm., in Newman and Ross, 1976:102), and *B. regalis*. *B. con*cavus concavus of Newman and Ross (1976) includes the two already known "B. c.indicus" and B. c. sinensis as synonyms of B. c. concavus (Newman and Ross, 1976). B. c. mexicanus lacking the transverse septa in the parietal tubes of shell wall does not belong to the *B. concavus* group in the strict sense, but is close to B. pacificus of the B. perforatus group pers. comm.). These four (Newman, extant members exclusive of B. c. mexicanus are living in Indo-Pacific, B. c. concavus in Dar-es-Salaam of Tanzania, the Persian Gulf and Hong Kong, and the other three species on the eastern Pacific. In the following paragraph, B. sendaicus is compared with these four "living" taxa of the *B. concavus* group on the basis of morphology.

*B. gregarius* (Conrad), 1856 was originally described as *Tamiosoma gregaria*, a new genus and species of a rudistid pelecypod by having the peculiar, elongate, vesiculose basis. This species is the most widespread and commonly encountered barnacle in the Miocene and Pliocene formations of central and southern California and Baja California regions (Zullo, 1964, 1969, 1979). "It was not until the opercular valves were described by Woodring", *in* Woodring *et al.* (1940), "that any

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Table 1. List of members of the B, concavus group and geographic and stratigraphic ranges quoted from Newman and Ross (1976) except for B. concavus concavus and "B. concavus indicus". [Oligo.: Oligocene, Mio.: Miocene, Plio.: Pliocene, Pleisto.: Pleistocene]

Balanus	bloxhamensis Weisbord, 1966:48, fossil, (Florida; Mio.)
В.	concavus concavus Bronn sensu de Alessanri, 1906:295, fossil (Europe; OligoPleisto.)
В.	c. alloplax Pilsbry and Olsson, 1951:200, fossil, (Ecuador; Oligo.)
Β.	c. chesapeakensis Pilsbry, 1916:103, fossil, (Maryland; Mio.)
В.	c. coosensis Dall, 1909:138, fossil, (Oregon; Mio.)
В.	c. dallonii Davadie-Suaudeau, 1952:20, fossil, (Algeria; Plio.)
В.	c. eseptatus Pilsbry, 1924:1, fossil, (Haiti; Mio.)
В.	c. finchii Lea, 1833:211, fossil, (Maryland; Mio.)
В.	c. glyptopoma Pilsbry, 1916:102, fossil, (east USA; Mio., Panama &
	east Mexico; Plio.)
В.	"c. indicus" Nilsson-Cantell, 1932:2, living, [Dar-es-Salaam, Tanzania]
В.	c. oligoseptatus Kolosváry, 1961:149, fossil, (USSR; Oligo.)
В.	c. proteus Conrad, 1834:134, fossil, (east USA; MioPlio.)
В.	c. raphanoides Moroni-Ruggieri, 1952:71, fossil, (Italy; Plio.)
В.	c. rariseptatus Pilsbry, 1918:186, fossil, (Panama; Mio.)
В.	c. rubescense Seguenza, 1876:450, fossil, (Italy; Tertiary)
В.	c. scutorum Seguenza, 1876:74, fossil, (Italy; Plio.)
В.	gregarius (Conrad), 1856:315, living, [California, 0-18 m], fossil, (central, southern &
	Baja California; MioPlio.)
В.	indicus Withers, 1923:291. fossil, (Pakistan; Mio.)
В.	panamensis Rogers, 1948:95, living, [Gulf of California to Pacific coast of Panama]
В.	polyporus Pilsbry, 1924:2, fossil, (Haiti; Mio.)
В.	regalis Pilsbry, 1916:108, living, [southern & Baja California]
В.	sendaicus Hatai, Masuda and Noda, 1976:13, fossil, (Japan; Mio.)
В.	talquinensis Weisbord, 1966:37, fossil, (Florida; Mio.)
В.	vadászi Kolosváry, 1949:2, fossil, (Europe; Mio.)

attempt could be made to determine the affinities of this unique species." "Woodriing concluded that *B. gregarius* was related to the *B. concavus* Bronn group." "Comparison of well preserved opercular valves with those of the extant Pacific coast species *B. aquila* reveals no apparent differences. The only major distinction in the two species would appear to be a tendency for the formation of elongate cellular basis in the fossils. However, this unique basis has now been discovered in extant specimens of *B. aquila* taken by the Allan Hancock Foundation off Santa Rosa Island, California. It is probable that the prevalence of elongate basis in fossils is related to past environmental conditions (warm, shallow embayments with high sedimentation rates) which are not widespread along the Pacific coast today." (Zullo, 1964:360). The morphological characteristics of this species, such as deeply sunken radii, growth ridge of scutum divided into beads by striations, and the rather narrow spur of tergum considerably close to the basiscutal angle,

Table 2. Morphological comparisons of B. sendaicus with the extinct B. concavus concavus sensu de Alessandri (1906) and the living taxa of the B. concavus group. ○: agreed; ×: disagreed. Informations on parietal tubes of shell wall in fossil European B. c. concavus and living B. gregarius are by courtesy of Newman (in press, pers. comm.).

	B. sendaicus	"B. concavus indicus" B. gregarius B. concavus concavus B. panamensis B. regalis
	surface smooth	$\circ \times \circ \circ \times$
/all	radii wide	$\bigcirc \times \times \times \times$
Shell w	summits of radii oblique	$\circ \times \circ \circ \times$
	parietal tubes with septa	×0000
Scutum	surface strongly sculptured	$\bigcirc \times \times \times \times$
	adductor ridge weak	$\bigcirc \times \times \bigcirc \times$
	spur furrow closed	×OO×O
Tergum	spur short	$\bigcirc \times \times \times \times$
	basal margin straight	× × 0 0 0

are different from those of *B. sendaicus. B. panamensis* Rogers, 1948 [Holotype, Calif. Acad. Sci. Dept. Paleo. Type Coll. 9434; paratypes, 9435-9437] was collected from C. A. S. Loc. 32532 of Chame Point, Canal Zone, Panama. Nature of shell form, toothed orifice, internal and basal margins of scutum, external surface of tergum, and basal margin of tergum found in this species are similar to those of *B. sendaicus*. Relatively narrow radii in comparison with *B. sendaicus*, weakly sculptured external surface of scutum, and shallow and opened spur furrow of tergum found in this species are apparently distinct from those of *B. sendaicus*. Nevertheless, among the American living taxa *B. panamensis* resembles the fossil Japanese *B. sendaicus*.

Β. regalis Pilsbry, 1916 [Holotype, USNM 43485] was described from Point Abreojos, west coast of Lower California based on shell wall without opercular valve. Rugged shell with irregularly arranged ribs of various size, extremely narrow radii, and moderately projecting and short adductor ridge of B. regalis distinguishes it from B. sendaicus. Morphological observation of opercular valves is based on samples obtained through the courtesy of Professor Newman. This species resembles *B. sendaicus* in the nature of spur furrow, parietal tubes of shell wall with transverse septa, and basal margin of tergum.

As mentioned above, the three living taxa distributed along the west coast of USA and central America are morphologically distinct from the fossil Japanese *B. sendaicus*. The morphological difference suggests the species level distinction between *B. sendaicus* and the three living taxa. Among the three living taxa, *B. panamensis* has more characters in common with *B. sendaicus* than the other two species. Thus it can be inferred that, among the living American species of the *B. concavus* group, *B. panamensis* is related to extinct *B. sendaicus*.

In Indo-Pacific area, survival of the infraspecific taxa of *B. concavus* is indicated by only three reports: Broch (1931), Nilsson-Cantell (1932) and Utinomi (1969). Their taxa were included in the *B. c. concavus* of Newman and Ross (1976). According to Newman and Ross (1976), studies on *B. c. concavus* except for the three reports mentioned above are concerned with fossil materials.

In Broch's type specimen of *B. c.* sinensis from Hong Kong, Newman and Southward "have distinguished two species, *B. reticulatus* and *B. variegatus*". "They have concluded that *B. c. sinensis* was a junior synonym of *B. variegatus* and that *B. reticulatus* of Utinomi remains valid (Newman and Southward, pers. comm.)" (Yamaguchi, 1980:1088).

Based on reexamination of Utinomi (1969)'s materials by Newman, *B. c. concavus* from the Persian Gulf reported by Utinomi (1969) is *Chirona* cf. *amaryllis* previously called *B. amaryllis*, because it does not have the parietal tubes in shell wall and has denticular ribs in inner base of parietes (Newman, pers. comm.). Therefore, the decision whether *B. c. concavus* of Newman and Ross (1976) is extant or not should be determined by reexamination of "*B. c. indicus*".

"B. concavus indicus" Nilsson-Cantell, 1932 was described based on only one specimen collected from a buoy in Dar-es-Salaam, Tanzania and has been preserved in the Berlin Museum. However, the type specimen can not be found in the Museum (Newman, pers. comm.). Live "B. c. indicus" has not been reported since 1932. Name indicus is preoccupied by B. indicus Withers, 1932. Therefore "B. c. indicus" is a homonym of B. indicus. However, both are apparently different from each other in morphology, so a replacement name is proposed by Newman (in press, pers comm.). Therefore, "B. c. indicus" of taxonomically invalid name is used in this paper.

Before entering to morphological comparison of *B. c. concavus* with "*B. c. indicus*", it will be historically reviewed on taxonomy of *B. c. concavus*. *B. concavus* was described by Bronn (1831) from the Pliocene of Andona Valley, near Asti, Piedmont, Italy (Pilsbry, 1916). However, the substance of *B. concavus* was not always clear until de Alessandri (1906)'s redescription of topotype materials.

Darwin (1854a) reported three different forms of terga under the name of B. concavus: 1) terga of pl. 1, figs. 4g from Coralline Crag, England (fossil) and of pl. 1, fig. 40 from Turin, Italy (fossil) having a narrow spur furrow and basal margin slightly depressed; 2) tergum of pl. 1, fig. 4k from Maryland, USA (fossil) having a very wide spur furrow and a straight basal margin; and 3) tergum of pl. 1, fig. 4m from Portugal (fossil, recent ?) having certainly opened spur furrow and basal margin slightly projected. These morphological differences are usually regarded to represent distinction of species level, therefore, it is not appropriate that those three forms of terga are summarized to a species of B. concavus. Darwin (1854b) also reported two different forms in terga under the name of B. concavus. Specimens of pl. 4, figs. 4d and 4e and specimen of pl. 4, fig. 4c in Darwin (1854b) are morphologically identical with those of pl. 1, figs. 4g and 4o and that of pl. 1, fig. 4m in Darwin (1854a) respectively.

After Darwin (1854a, b), studies of B. concavus have been subjected to his great influence, and three distinct forms have probably been treated as a species of B. concavus. This fact seems to be a main cause for the taxonomic confusion of B. concavus.

Revision of topotypes (pl. 16, figs. 21– 25; pl. 17, figs. 1a-4b) of *B. concavus* Bronn by de Alessandri (1906) seems to be found a clue to solve the taxonomic confusion and pointed out that terga of the real *B. concavus* are identical with those of pl. 1, figs. 4g and 40 in Darwin (1854a) and of pl. 4, figs. 4d and 4e in Darwin (1854b). Therefore, *B. concavus* redescribed by de Alessandri (1906) is called *B. c. concavus* sensu de Alessandri (1906) in this paper. It is interpreted that B. c. concavus of Newman and Ross (1976) contains five or more species of B. c. concavus sensu de Alessandri, "B. c. indicus", B. c. sinensis (=B. amphitrite s. 1.), and at least two species of Balanus (?) eliminated B. c. concavus sensu de Alessandri from Darwin (1854a)'s B. concavus.

In this paper, *B. concavus* in the following references is considered to be identical with *B. concavus concavus* Bronn in the sense of de Alessandri (1906): Darwin, 1854a, p. 17 (in part), figs. 4g and 4o; Darwin, 1854b, p. 235 (in part), pl. 4, figs. 4d and 4e; de Alessandri, 1895, p. 282, pl. 2, figs. 10a-10e; Pilsbry, 1916, p. 100, pl. 21, figs. 1-1c; Davadie, 1963, p. 52, pl. 24, figs. 8-9, pl. 26, figs. 1-4, pl. 28, figs. 1a-2b; and Menesini, 1965, p. 110, pl. 28, figs. 2-6, pl. 29, figs. 1-9, pl. 30, figs. 1-4, pl. 34, figs. 7-8, pl. 35, figs. 1-2, pl. 41, figs. 3-8, pl. 42, figs. 1-8, pl. 43, figs. 1-3.

"B. c. indicus" and B. c. concavus sensu de Alessandri (1906) have common morphological characteristics, such as conical shell form, rhomboidal orifice, oblique summits of radii, external scutal surface with longitudinal striations, relative spur width for basal length of tergum, relative position of spur, shape of basal margin of tergum, etc. However, the following morphological differences are present: shell surface is smooth in "B. c. indicus", but costated, sulcated, festooned, or with longitudinal ribs in B. c. concavus. Parietal tubes of shell wall have not transverse septa in "B. c. indicus", but have transverse septa in B. c. concavus (Information on this shell structure is by courtesy of Professor Newman, pers. comm.). The adductor ridge is very low and short in "B. c. indicus", but prominent, long, clear and slightly curved in B. c. concavus. Tergal surface has longitudinal striae on carinal side of "B. c. indicus", but is longitudinally and transversally sulcated in *B. c. concavus*. Spur furrow is open in "*B. c. indicus*", but closed in *B. c. concavus*. The morphological resemblances show close relationship between them, however, the morphological distinctions are usually sufficient to recognize them as different species. As mentioned above, *B. c. concavus* sensu de Alessandri (1906) is neither an extant species nor "*B. c. indicus*", therefore, *B. c. concavus* is an extinct species.

B. sendaicus and B. c. concavus sensu de Alessandri (1906) closely resemble each other in their morphology. The common characteristics are conical shell shape, rhomboidal and toothed orifice, having transverse septa in parietal tubes of shell wall, striated external surface of scutum, etc. In B. sendaicus, shell wall is generally smooth, but occasionally costated in carina and carinolateral. This character resembles that of B. c. concavus. The two species, however, differ morphologically in the following characters: high, long and clear adductor ridge of scutum in B. c. concavus, but obscure in B. sendaicus; long spur length in B. c. concavus, but moderately long in B. sendaicus; and surface of tergum longitudinally, and transverselly sulcated in B. c. concavus, but having only weak growth lines in B. sendaicus. These morphological differences are sufficient enough to distinguish them as two distinct species. Morphological resemblance, however, suggests their close relationship.

Furthermore, when *B. sendaicus* and "*B. c. indicus*" compared on the basis of morphology, they are found to resemble each other. The conical form, rhomboidal and moderately wide orifice, smooth shell surface (in *B. sendaicus*, carina and carinolateral are often costated, but generally smooth), scutum with well developed growth lines and strong striae on external

surface, and crenulation of basal margin of scutum are their common character. However, they have four different characteristics: parietal tubes of shell wall have not transverse septa in "B. c. indicus", but have the septa at upper half of shell wall in B. sendaicus, though usually secondarily filled up with calcite as result of fossilization; the adductor ridge of scutum is very low and short in "B. c. indicus", but obscure in B. sendaicus; distance from basiscutal angle to spur of tergum approximately is equal to spur width in "B. c. indicus" (Henry and McLaughlin, 1975), but is 2/3 in B. sendaicus (average 25 individuals examined); and the spur furrow is open in "B. c. indicus", but partially closed in B. sendaicus. Thus, morphological distinction is clear and sufficient enough to separate them into two distinct species.

*B. sendaicus* was described by Hatai *et al.* (1976) based on only one scutum (Holotype). However, the shell walls of *B. sendaicus* were described under the name of *B. rostratus* Hoek, and the tergum was undescribed. Thus, *B. sendaicus* of Hatai *et al.* (1976) clearly does not agree with *B. sendaicus* redescribed here. However, as mentioned above, *B. sendaicus* redefined here is not identical with the already known living species of the *B. concavus* group.

# Comparison with the extinct taxa of the *B. concavus* group

*B. bloxhamensis* Weisbord, 1966 [Holotype, PRI (Paleontological Research Institution) 27408] was described from the late Miocene Choctawhatchee Formation at Jackson Bluff in Leon County, Florida. This species is based on complete shells without opercular valve, and "the V-shaped cut at the summit of the radii, the longitudinal slit dividing the radii, and the vertical ridges transversing the radii, are the distinctive characters of the shell". These characters are distinct from those of *B. sendaicus*.

B. concavus alloplax Pilsbry and Olsson, 1951 [Holotype, ANSP (Academy of Natural Sciences of Philadelphia) 18869] is represented only by opercular valves from the Oligocene age of Playa Rica, Rio Santiago, Ecuador. Shell walls without opercular valves were described from another fossil locality, Posorja, Ecuador. It is difficult whether these shell walls belong to the same subspecies, until more fossil specimens from the type locality The following morphoare available. logical characteristics of this subspecies are distinct from those of B. sendaicus; scutum strongly wrinkled; tergum with "a decidedly sunken spur fasciole"; spur removed "from basiscutal angle by nearly twice width of spur".

B. concavus chesapeakensis Pilsbry, 1916 [Holotype, ANSP 1143] is based on a complete shell having scutum and tergum from the Miocene beds of Maryland. A long and high adductor ridge in its middle part, and lateral depressor muscle pit having a strongly prominent bordering ridge, parallel to adductor ridge in B. c. chesapeakensis distinguish it from B. sendaicus.

*B. concavus coosensis* Dall, 1909 [Holotype, USNM 153960] was originally described as a variety of *B. tintinnabulum*, based on four specimens of shell wall without opercular valves from the Miocene of Coos Bay, Oregon. As Pilsbry (1916) poined out, *B. concavus coosensis* having oblique summit radii and no porous radii is not related to *B. tintinnabulum*. Very large specimen attaining 50 mm in greatest diameter and the relatively narrow rostrum typical of *B. concavus coosensis* is not found in *B. sendaicus*.

B. concavus dallonii Davadie-Suaudeau,

1952 is based on a specimen of shell wall from the Astian (late Pliocene) of Aïr Lalla Barta, Algeria. The shell surface having ribs and striations and large orifice distinguishes it from those of *B. sendaicus*.

*B. concavus eseptatus* Pilsbry, 1924 [Holotype, USNM 352257] was described based on complete shell wall without opercular valve from the Miocene of Haiti. In *B. concavus eseptatus*, the radii are narrower and the orifice is larger than those of *B. sendaicus*.

*B. concavus finchii* Lea, 1833 was originally described as *B. finchii* Lea from the Miocene of St. Marys, Maryland. Pilsbry (1930: 432) redescribed and illustrated topotype specimens. This subspecies is distinguished from *B. sendaicus* by large orifice, weak growth ridges and a long, moderately well developed adductor ridge in scutum.

*B. concavus glyptopoma* Pilsbry, 1916 [Syntype, ANSP 1140] was described from the Pliocene Caloosahatchie Formation, Florida. The rather numerous small ribs, sometimes weak or nearly obsolete on the surface of parietes, a strong but rather short adductor ridge, and a short, widely obliquely truncated spur of this subspecies distinguish it from *B. sendaicus*.

*B. concavus oligoseptatus* Kolosváry, 1961 [Holotype preserved in Coll. Mus. Inst. Syst. Zool. Univ. Szeged] from the upper Oligocene of Kizil-Kum, USSR has level summits radii and large orifice, which are different characteristics from *B. sendaicus*.

*B. concavus proteus* Conrad, 1834 was described based on specimens from James River, Suffolk, and Yorktown, Virginia. This subspecies is characterized by strong ribs on the external shell wall, and slightly well developed adductor ridge. These morphological characteristics are not present in *B. sendaicus*.

*B. concavus raphanoides* Moroni-Ruggieri, 1952 [Holotype, Cirr. 342] from the Pliocene of Bagni of Casciana, Italy is represented by an internal mould of shell wall. The characteristic cup-shaped base, cylindrical form and wide orifice of this subspecies are not present in *B. sendaicus*.

*B. concavus rariseptatus* Pilsbry, 1918 [Holotype, USNM 324292] is based on one cluster of shell walls without opercular valves from the Miocene Gatum Formation of Panama Canal Zone. The slightly longitudinally roughened walls of characteristic of this subspecies are not present in *B. sendaicus*.

B. concavus rubescens Seguenza, 1874 was originally described as a variety of B. concavus from the Tertiary formation of the province of Messina, Italy. In B. concavus rubescens, the growth ridges of scutum are scarcely detectable and the longitudinal striations on external surface of scutum are weak. These morphological characteristics are distinct from those of B. sendaicus.

*B. concavus scutorum* Seguenza, 1876 was described as a distinct species *B. scutorum* from the province of Messina, Italy. The following morphological characteristics of this subspecies are different from those of *B. sendaicus*: surface of scutum with fine longitudinal striations; spur width nearly equal to 1/3 basal length of tergum; narrow spur furrow; distinctly prominent adductor ridge.

*B. indicus* Withers, 1923 [Holotype, In. 20233] was reported from the Miocene of Karachi, Mouth of the Indus. This species is distinguished from *B. sendaicus* by rather deeply sunken radii, narrowly spaced and weakly prominent growth ridges on scutum, and spur fasciole in tergum.

*B. polyporus* Pilsbry, 1924 [Holotype, USNM 352258] from the Miocene of Haiti is represented by only one shell wall specimen. The somewhat roughened wall surface and rostrum notably longer than the other plates distinguish it from *B*. sendaicus.

*B. talquinensis* Weisbord, 1966 [Holotype, PRI 27374] was described from the Miocene Choctawhatchee Formation at Jackson Bluff in Leon County, Florida. The following morphological characteristics of *B. talquinensis* are distinct from those of *B. sendaicus*: small beads or nodules are found at the intercepts of the radial riblets with the growth ridges of scutum; adductor ridge of scutum is high and long; spur furrow of tergum is shallow and moderately wide.

*B. vadászi* Kolosváry, 1949 was reported from the early Miocene of Felsötarkány, Hungary. Compartments with narrow radii and coarse white ridges on surface and scutum without longitudinal striations on external surface of *B. vadászi* are not found in *B. sendaicus*.

In conclusion, as mentioned above, *B.* sendaicus is not identical with any of the already known living and extinct taxa of the *B. concavus* group, and is surely a distinct species as pointed by Hatai *et al.* (1976). However, because *B. sendaicus*  based on a single scutum of Hatai *et al.* (1976) is not entirely identical with that of this paper based on complete specimens and description given by Hatai *et al.* (1976) is too brief, *B. sendaicus* is redescribed here.

# Systematic Paleontology

Superfamily Balanoidea Leach, 1817 Family Balanidae Leach, 1817 Genus *Balanus* Da Costa, 1778

Balanus sendaicus Hatai, Masuda and Noda, 1976

- [Pl. 44, Figs. 1-6, Pl. 45, Figs. 1-9,
- Pl. 46, Figs. 1-22, Pl. 47, Figs. 1-22, Pl. 48, Figs. 1-3, Text-fig. 1]
- Balanus concavus Bronn: Yamaguchi, 1974, p. 218, pl. 67, figs. 9-10.
- Balanus rostratus Hoek: Hatai, Masuda and Noda, 1976, p. 14, pl. 2, figs. 1, 2, 13-17, 19-21.
- Balanus sendaicus Hatai, Masuda and Noda, 1976, p. 13, pl. 2, figs. 1 (same specimen as fig. 20), 9 (same specimen as fig. 10, Holotype, missing), 10, 14-17, 19, and 20.

Explanation of Plate 44

Balanus sendaicus Hatai, Masuda and Noda, 1976 from the Miocene Moniwa Formation

All specimens were illustrated under the name of *Balanus rostratus* Hoek in plate 2 of Hatai *et al.* (1976) and were deposited in the Saito Ho-on Kai Museum, Sendai, Japan.

[Hatai, et al.'s specimens are not accompanied with registration number.]

- Fig. 1. la-b, shell wall (same specimen as pl. 2, figs. 1 and 20 of Hatai et al.) on gravel, ×1.
- Fig. 2. 2a-c, shell walls (same specimens as pl. 2, fig. 2 of Hatai *et al.*) on a brachiopod shell of *Coptothyris grayi miyagiensis* Hatai, Masuda and Noda, 2a, ×1. 2b-c, ×2. 2d-k, opercular valves preserved in the shell cavity of shall (left) [These opercular valves were not found by Hatai *et al.*), ×5; 2d-e, right scutum. 2f-g, right tergum. 2h-i, left scutum [morphologically identical with Holotype (missing), pl. 2, figs. 9 and 10 of Hatai *et al.*], 2j-k, left tergum.
- Fig. 3. 3a-b, shell wall (same specimen as pl. 2, fig. 14 of Hatai *et al.*) on a brachiopod shell of *Coptothyris grayi miyagiensis*, 3a, ×1. 3b, ×2.
- Fig. 4. 4a-b, shell wall (same specimen as pl. 2, fig. 15 of Hatai *et al.*) on a brachiopod shell of *Coptothyris grayi miyagiensis*, 4a, ×1. 4b, ×2.
- Fig. 5. 5a-b, shell walls (same specimens as pl. 2, fig. 19 of Hatai *et al.*) on a molluscan shell of *Anomia* sp., 5a-b,  $\times 1$ .
- Fig. 6. 6a-b, shell walls (same specimens as pl. 2, figs. 16 and 17 of Hatai et al.) on gravel, ×1.





*Description.*—General appearance.-Shell conical and smooth often with weak longitudinal ribs. Color of shell ivory yellow, frequently with very fine striations of pale reddish purple or tinted with pale reddish purple near orifice or on the shell wall. Orifice toothed.

Compartments.-Parietes very wide in rostrum and lateral. Radii wide in rostrum and lateral, and moderately wide in carinolateral and carina; summits oblique. Sheath solid. Parietal (longitudinal) tubes with transverse septa in upper half of shell wall. Parietal tubes usually secondarily filled up with calcite as the result of fossilization. Those transverse septa usually not found by soft X radiographs (Pl. 48, Figs. 15-19), although, they observable in the longitudinal thin section of shell wall (Text-fig. 1, Pl. 48, Figs. 1-3).

Base.-Calcareous.

Opercular plates.-Scutum slightly concave externally with narrowly inflected tergal segment; basal angle obliquely truncated; external surface with prominent and crenulate growth ridges and moderately prominent longitudinal ridges frequently bifurcated by the weak longitudinal striae; nodes observable at the connection between the ridges; these ridges circumscribing square depressions ("rather coarsely latticed sculpture" as pointed by Pilsbry, 1916:101); color ivory yellow; inner surface smooth on upper part; articular ridge moderately projected, reflexed and greater than 1/2 length of tergal margin; articular furrow narrow and shallow; adductor ridge weak or hardly visible; pit for adductor muscle wide and deep; pit for lateral depressor muscle wide and deep; inner surface of basal margin faintly crenulated.

Tergum flat, moderately wide with a weakly developed beak (Pl. 44, Figs. 2f-g); external surface smooth with fine growth lines; color ivory yellow, tinted



Text-fig. 1. Shell structures in longitudinal sections of a rostrum (UMUT CA9686) from the Miocene Moniwa Formation (Loc. 4). Photos for Text-figs. 1a-b are shown in Pl. 48, Figs. 1-2. 1a, a parietal tube with transverse septa in a longitudinal section parallel to the parietal (longitudinal) tube. 1b, parietal tubes with transverse septa in an oblique longitudinal section. I: inner lamina, Lr: longitudinal rib, Ls: longitudinal septum, O: outer lamina, P: parietal tube, S: sheath, T: transverse septum.

with pale reddish purple near apex between spur furrow and carinal margin; basal margin of both sides of spur approximately straight; spur furrow extremely narrow in young specimens or partially closed in large specimens and slightly sunken; spur rounded at end and with several very weak projections on inner surface of basal margin; spur width approximately 1/3 length of basal margin or less; distance between basiscutal angle and anterior side of spur nearly equal to 1/3 length of spur width; crests for depressor muscle well developed; rest of inner surface smooth.

# Stratigraphic and geographic ranges, definition of *B. c. concavus* and habitat

The straigraphic range of each member of the B. concavus group summarized by Newman and Ross, 1976 is shown in Table 1. The oldest record is the Oligocene of Nurri, Sardinia Is. and Capo S. Elia, Cagliari, Italy concerning B. c. concavus (de Alessandri, 1906), of Pomaz and God, Hungary concerning B. c. concavus (Kolosváry, 1943), of Ecuador concerning B. c. allopax (Kolosváry, 1943), of Budafork concerning B. c. chesapeakensis and B. c. glyptopoma (Kolosváry, 1943), and of Kizil-Kum, USSR concerning B. c. (Kolosváry, oligoseptatus 1943). The youngest record is the Pleistocene of Monte Mario, Roma; Civitavecchia, Palo; and Sciacca, Sicilia Is., Italy concerning *B. c. concavus* (de Alessandri, 1895, 1906), of Pisa province, Italy concerning *B. c. concavus* (Menesini, 1965), and of Monastir, Tunisia concerning *B. c. concavus* (Davadie 1963), etc. However, most of the fossils were obtained from the Miocene and Pliocene.

There are two critical and important problems of stratigraphic distribution of the *B. concavus* group:

1) Most of *B. c. concavus* reported in the literature appear only as listed occurrence of fossil and/or have only brief morphological description without illustration. Are most of these specimens identical with *B. c. concavus* in the sense of de Alessandri (1906, pl. 16, figs. 21-25; pl. 17, figs. 1a-4b), based on reexamination of topotypes of Bronn (1831)'s *B. concavus*? The writer discussed precisely on *B. concavus* of Darwin (1854a, b) and of de Alessandri (1906) in the preceding line.

2) Geologic age-determination in Neogene and Quaternary were advanced by biostratigraphy based on planktonic microfossils since 1950's and by magnetostra-

# Explanation of Plate 45

#### Balanus sendaicus Hatai, Masuda and Noda, 1976

All specimens are topotypes from Loc. 4 of the early to middle Miocene Moniwa Formation and preserved in University Museum, University of Tokyo (UMUT).

- Fig. 1. a complete shell (UMUT CA9641). 1a-b, a shell wall (conical form), ×1. 1c, four opercular valves articulated, ×2.
- Fig. 2. a complete shell (UMUT CA9642). 2a-b, a shell wall (cylindrical form), ×1.5. 2c-j, four opercular valves, ×2.5; 2c-d, right scutum. 2e-f, left scutum. 2g-h, right tergum. 2i-j, left tergum.
- Fig. 3. a complete shell (UMUT CA9643). 3a-c, a shell wall (cylindrical form), ×1.5. 3d-k, four opercular valves, ×2.5; 3d-e, right scutum. 3f-g, left scutum. 3h-i, right tergum. 3j-k, left tergum.

Fig. 4. a shell wall (UMUT CA9644). 4a-b, ×1.5 (a single right scutum preserved in shell cavity).

- Fig. 5. an aggregation of shell walls (UMUT CA9645),  $\times 1.5$ .
- Fig. 6. shell walls (UMUT CA9646). 6a-b,  $\times 1.5$ .
- Fig. 7. a shell wall (UMUT CA9688) on fragment of brachiopod shell. 7a-b, ×1.5.
- Fig. 8. a shell wall (UMUT CA9647). 8a-b,  $\times 1$ .
- Fig. 9. a shell wall (UMUT CA9648). 9a-b,  $\times 1$ .

Plate 45



tigraphy based on paleomagnetic polarity in sediments since 1970's. Determination of the stratigraphic range of *B. concavus* requires knowledge of the precise geologic age of the sampling horizons. However, many previous studies of *B. concavus* did not incorporate the current stratigraphic method mentioned above in age-determination, because they were published before the recent advance of age-determination.

Stratigraphy of Italy, based on modern geologic views, was established after the World War II and thus the geologic ages published before the World War II will have to be compared with the advanced stratigraphy and geologic age.

Therefore, in case where these two important problems of identication or definition of species and geologic age are not concurrently solved, discussion on stratigraphic and evolutionary change of *B. c. concavus* or other members of the *B. concavus* group should be restricted until the precise range of the *B. concavus* group in Mediterranean region is summarized.

The fossils of this group were found in areas surrounding the Mediterranean. corresponding to the western part of Tethys Sea (Oligocene-Pleistocene), England (Pliocene), west Atlantic coast of North and South America (Pliocene), the east Pacific coast from California to Mexico (Miocene-Pliocene), and east Pacific coast of Peru (Pleistocene). One noteworthy result is, as pointed by Pilsbry (1916), that this group had flourished from the Miocene to Pliocene and had lived in the Atlantic and the Mediterranean region until Pleistocene, but is extinct in this region ever since, though its members continuously live on the east Pacific coast of California and Baja Cailfornia at present. The disappearance of this group from the Atlantic and the Mediterranean

region in Pleistocene is a mystery.

In the east Pacific region, after the closure of Panama isthmus in the Miocene, three species: *B. gregarius* (=*B. aquila*, a junior synonym of *B. gregarius*), *B. panamensis*, and *B. regalis*, live in the California and Oregonian Provinces from California to Peru-Chile. They are relicts of the Tethys (Newman, 1979). *B. gregarius* is endemic to the narrow Transition Zone (about 4° in latitude) in between California and Oregonian Provinces (Newman, 1979). The nature of this "so-called 4° short-range endemism" is discussed in detail by Newman (1979).

Japanese fossil *B. sendaicus* is the first record of this species group from the western Pacific area. *B. sendaicus* were collected from several localities of northern Honshu, Japan. The geologic age is restricted to the Miocene.

B. sendaicus occurs in association with extinct pectinid Chlamys kaneharai, brachiopods and bryozoans from the type locality of this species (Loc. 4). Brachiopods are well preserved, and the majority of them are articulated. Bryozoans have fine and reticulated structure, and are also well preserved. Shells of B. sendaicus are chiefly found as fragmental, skeletal elements, however, some of shells are complete shell walls and in some cases opercular valves are preserved in their shell cavities. These facts seem to indicate that those fossils were not transported so far away from their habitat. Those fossils associated indicate the paleoenvironment of moderately strong water agitation. The habitat of B. sendaicus is also presumed from the nature of fossil assemblage and sediments at the other fossil localities of B. sendaicus.

# Conclusion

The description of Balanus sendaicus

Hatai, Masuda and Noda (1976) was based on a single scutum, but it could not found out in the Saito Ho-on Kai Museum, Sendai. Study of topotype materials suggests that *B. sendaicus* is conspecific with shell walls incorrectly identified as *B. rostratus* Hoek by Hatai *et al.* (1976). *B. sendaicus*, based on scuta, terga and shell walls, is distinct from the known taxa, and is a good species as proposed by by Hatai *et al.* (1976). *B. sendaicus* is redescribed on the basis of skeletal elements of the topotypes.

*B. sendaicus* belongs to the *B. concavus* group based on having transverse septa in parietal tubes of shell wall and the characteristic ornamentation of scutum and seems to have derived from the closely related *B. concavus concavus* Bronn sensu de Alessandri (1906) which has mainly lived in the Mediterranean region of Italy and its adjacent areas from the Oligocene to the Pleistocene.

Darwin (1854a, b)'s *B. concavus* includes two or more taxa which are morphologically distinct from true *B. concavus concavus*. Newman (in press) reexamined based on the materials of Darwin (1854a, b), and recognized three taxa including true *B. concavus concavus* (Newman, pers. comm.). It is difficult to decide whether or not extinct infraspecific taxa of *B. concavus* are distinct from the other taxa, judging from their brief descriptions and their incomplete assemblage of skeletal elements. Their taxonomic positions will be valid by the Zoological Nomenclature, however, their holotypes and topotypes are necessary to be reexamined from the modern taxonomic point of view.

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# Explanation of Plate 46

Balanus sendaicus Hatai, Masuda and Noda, 1976

All specimens are topotypes from Loc. 4 of the early to middld Miocene Moniwa Formation and preserved in University Museum, University of Tokyo (UMUT).

- Figs. 1-4. rostrum, ×1.2. 1a-b, (UMUT CA9649). 2a-b, (UMUT CA9650). 3a-b, (UMUT CA9651). 4a-b, (UMUT CA9652).
- Figs. 5-8. lateral, ×1.2. 5a-b, left, (UMUT CA9653). 6a-b, left, (UMUT CA9654). 7a-b, left, (UMUT CA9655). 8a-b, right, (UMUT CA9656).

Figs. 9-12. carinolateral, ×1.2. 9a-b, left, (UMUT CA9657). 10a-b, left, (UMUT CA9658). 11a-b, left, (UMUT CA9659). 12a-b, right, (UMUT CA9660).

Figs. 13-16. carina, ×1.2. 13a-b, (UMUT CA9661). 14a-b, (UMUT CA9662). 15a-b, (UMUT CA9663). 16a-b, (UMUT CA9664).

Figs. 17-19. scuta. 17a-b, right, (UMUT CA9665), ×3. 18a-b, left, (UMUT CA9666), ×2.5. 19a-b, left, (UMUT CA9667), ×2.5.

Figs. 20-22. terga, ×2.5. 20a-b, right, (UMUT CA9668). 21a-b, right, (UMUT CA9669). 22a-b, left, (UMUT CA9670).

YAMAGUCHI: Miocene Cirriped Balanus sendaicus





Text-fig. 2. Sampling localities of fossil *Balanus sendaicus*. Details for each site are given in the section Locality.

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# Locality

la-b.—Obuchi: a) cliff (40°58′55.1″N, 141° 21′42.9″E) along the Oippe river, situated about 2.85 km N26°W of Obuchi, Rokkashomura, Kamikita-gun, Aomori Pref., crosslaminated coarse shell sand ("coquina")., b) cliff (40°58′55.1″N, 141°20°49.3″E) along the Oippe river, situated about 3.55 km N44°W of Obuchi; cross-laminated coarse shell sandstone ("conquina")., a-b) late Miocene Takahoko Formation.

2a-c.—Ohmagari: a) shell quarry at Arasawa (39°24′16.2″N. 140°21′4.2″E), Minami-Sotomura, Senpoku-gun, Akita Pref.; barnacle shell sand ("coquina")., b) cliff (39°21'34.1"N, 140°21'42.1"E) at Asizawa Pass of Takinoue, Ohmori-machi, Hiraga-gun, Akita Pref.; massive poorly sorted tuffaceous sandstone containing carbonaceous matter., c) cliff (39° 22'6.5"N, 140°21'42.1"E) of left bank of creak situated about 1 km north of Loc. 2b; massive poorly sorted tuffaceous fine sandstone containing carbonaceous matter., a-c) late Miocene Sugota Formation.

3.—Ukibuta: cliff (39°18'47.8"N, 140°20'41.1" E) situated about 100 m east of Ukibuta Tunnel, Ukibuta, Higashi-Yuri-mura, Yuri-gun, Akita Pref.; poorly sorted, tuffaceous, coarse sandstone; late Miocene Sugota Formation.

4.—Sendai: cliff (type locality of *B. sendaicus*) (38°12'45.4"N, 140°47'26.9"E) situated about 200 m south of the Oide Bridge across the Natori river, Moniwa, Sendai City, Miyagi Pref.; poorly sorted, gravelly, calcareous, medium, shell sandstone; early to middle Miocene Moniwa Formation.

5.—Nanao: cliff (37°2′23.5″N, 136°57′23.4″E) situated about 0.95 km N85°W of Nanao Station, Iwaya-cho, Nanao City, Ishikawa Pref.; cross-laminated conglomeratic coarse shell sandstone containing many bryozoans; late Miocene Nanao Calcareous Sandstone.

6a-b.—Mizunami: a) cliffs (35°23'32.5N", 137°16'35.6"E) situated about 3.5 km N32°E of Mizunami Station, Dan, Mizunami City, Gifu Pref.; conglomeratic very coarse sandstone., b) cliff (35°22'25.7"N, 137°16'28.6E) situated about 1.95 km N61°E of Mizunami Station, Sakurado, Mizunami City., a-b) late Miocene Oidawara Formation.

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Explanation of Plate 47

#### Balanus sendaicus Hatai, Masuda and Noda, 1976

All specimens are from various localities of the Miocene and preserved in University Museum, University of Tokyo (UMUT).

- Fig. 1. Specimen from the late Miocene Takahoko Formation (Loc. 1). 1a-b, left scutum (UMUT CA9671), ×2.5.
- Fig. 2. Specimen from the late Miocene Sugota Formation (Loc. 2a). 2a-b, left scutum (UMUT CA9672), ×2.5.
- Figs. 3-5. Specimens from the late Miocene Sugota Formation (Loc. 2b). 3a-b, left scutum (UMUT CA9673), ×4. 4a-b, left scutum (UMUT CA9674), ×4. 5a-b, right scutum (UMUT CA9675), ×4.
- Figs. 6-10. Specimens from the late Miocene Sugota Formation (Loc. 2c). 6a-b, left scutum (UMUT CA9676), ×3. 7a-b, right scutum (UMUT CA9677), ×4. 8a-b, left tergum (UMUT CA9678), ×4. 9a-b, left tergum (UMUT CA9679), ×4. 10a-b, right tergum (UMUT CA9680), ×4.
- Figs. 11, 12. Specimens from the late Miocene Nanao Calcareous Sandstone (Loc. 5). 11a-b, right scutum (UMUT CA9681), ×4. 12a-b, left scutum (UMUT CA9682), ×4.
- Figs. 13, 14. Specimens from the late Miocene Oidawara Formation (Loc. 6b). Coll. Mizunami Fossil Museum, Mizunami, Gifu Pref. 13a-b, left scutum (UMUT CA9683), ×2.5. 14a-b, right scutum (UMUT CA9684), ×3.
- Figs. 15-19. X-radiographs of topotype specimens from the Miocene Moniwa Formation (Loc. 4), ×1. 15, carina (UMUT CA9661). 16, rostrum (UMUT CA9649). 17, rostrum (UMUT CA9651).
  18, rostrum (UMUT CA9652). 19, left lateral (UMUT CA9653).
- Figs. 20-22. Interlaminate figures observed in thin section parallel to the base of shell wall (Loc. 4). 20, carina (UMUT CA9685). 21, rostrum (UMUT CA9686). 22, left lateral (UMUT CA9687).



Plate 47



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# Explanation of Plate 48

Shell structure in longitudinal sections of a rostrum (UMUT CA9686) from the Miocene Moniwa Formation (Loc. 4).

- Fig. 1. a parietal tube with transverse septa in a longitudinal section parallel to the parietal (longitudinal) tube. 1a, ×8.4. 1b, ×10.2. 1c, ×42.
- Fig. 2. parietal (longitudinal) tubes with transverse septa in an oblique longitudial section (four parietal tubes are found). 2a, ×5.8. 2b, ×10.2. 2c, ×42.
- Fig. 3. parietal (longitudinal) tubes with transverse septa in an oblique longitudinal section. 3a, photos by an optical microscope of Replica film (cellulose acetate film),  $\times 10.2$ . 3b-e, external and internal walls (correspond to the outer and inner laminae) of parietal tubes and transverse septa partially left the original shell structure. Interspace of the parietal tube is secondarily filled up with the calcite which contrasts with the calcite composing the transverse septa in grain size, shape and arrangement. Photos by scanning electron microscope (SEM) of the etched surface. 3b, enlargement of top of fig. 3a,  $\times 42$ . 3c, ditto,  $\times 128$ . 3d, ditto,  $\times 290$ . 3e, ditto,  $\times 728$ .



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Obuchi 尾駮, Oippe 老部, Takahoko 鷹架, Ohmagari 大曲, Arasawa 荒沢, Asizawa 芦沢, Takinoue 滝ノ上, Sugota 須郷田, Ukibuta 浮蓋, Sendai 仙台, Oide 生出, Natori 名取, Moniwa 茂庭, Nanao 七尾, Mizunami 瑞浪, Dan 段, Sakurado 桜堂, Oidawara 生俵

日本の中新世蔓脚類 Balanus sendaicus Hatai, Masuda and Noda について----テー チス海の Balanus concavus 種群と比較して----: Balanus sendaicus は仙台付近に分布す る中新世茂庭層から得られた殻の一部分の1つの楯板 (scutum) に基づいて畑井,増田,野田 (1976) により新種として記載された。その標本は仙台の斉藤報恩会博物館に保管されている はずだが、そこには見つけられなかった。その名前のもとに殻の他の部分、背板 (tergum) お よび周殻 (shell wall), は記載されなかった。原産地標本 (topotype) および畑井ら (1976) に図示された標本調査で、畑井らが B. rostratus として記載・図示した周殻は B. sendaicus に属す。それゆえ、本論文では畑井らによって記載されなかった背板を含めて B. sendaicus を明確にする。B. sendaicus は畑井らの指摘のとうり既知の分類群から異なる。B. sendaicus は漸新世から更新世に地中海地域テーチス海に繁栄した de Alessandri (1906)の意味合いで の絶滅種 B. concavus と近縁である。B. sendaicus は周殻の壁管 (longitudinal または parietal tubes) に直交副隔壁 (transverse septa) を持つことおよび楯板の特徴的な表面装 飾によって B. concavus に代表される種群に属す。この種群に属す"4 種"が東部太平洋に 生息している (Newman, 1979) けれども, 大西洋・地中海地域からの B. concavus 種群 の消滅は Pilsbry (1916) の指摘のようにミステリアスである。化石の日本の B. sendaicus は、中新世に限られ、B. concavus 種群の東アジアからの最初の化石記録となり、そしてテ ーチス海と東アジアの間のつながりの存在を指摘する。 山口寿之

**Postscript:** After submitted, Saito Ho-on Kai Museum of Natural History (SHM) let me know on registration numbers of illustrated specimens in the plate (pl. 2) of Hatai *et al.* (1976). The registration numbers are as follows.

Registration number	figure in Hatai <i>et al</i> .	species name in Hatai <i>et al.</i>	species name in this paper
SHM, Reg. No. 22069	figs. 1 & 20	Balanus rostratus	B. sendaicus
22070	figs. 3 & 4	B. oidensis (Holotype)	Not studied
22071	figs. 5 & 6	B. moniwaensis (Ditto)	Not studied
22072	figs. 7 & 8	B. natorianus (Ditto)	Not studied
22073	figs. 11 & 22	B. miyagiensis (Ditto)	Not studied
22074	fig. 12	B. tohokuensis (Ditto)	Not studied
22075	fig. 18	B. rikuzenensis (Ditto)	Not studied
22076	fig. 23	<i>B</i> . sp.	Not studied
22077	fig. 2	B. rostratus	B. sendaicus
22078	figs. 13 & 21	B. rostratus	B. crenatus
22079	fig. 14	B. rostratus	B. sendaicus
22080	fig. 15	B. rostratus	B. sendaicus
22081	figs. 16 & 17	B. rostratus	B. sendaicus
22082	fig. 19	B. rostratus	B. sendaicus
22083	figs. 9 & 10	B. sendaicus (Holotype, missing)	B. sendaicus

# Palaeontological Society of Japan Special Papers No. 24 Permian Conodont Biostratigraphy in Japan

By Hisaharu IGO

Issued December 15, 1981, 50 pp., 12 pls.

Price (postage and handling incl.) U.S. \$24.00.

Clark and Behenken (1971) summarized "Conodont and biostratigraphy of the Permian" in Symposium on Conodont Biostratigraphy edited by Sweet and Bergström. In this paper, they mentioned that Permian System represents the "last frontier" for conodont research. Permian conodont study was the same situation in Japan.

About ten years ago, Hisaharu Igo initiated to study Permian conodonts and he succeeded to recover numerous specimens from the various sections which were dated by fusulinaceans mainly in central part of Japan. The author clarified stratigraphic as well as geographic distribution of many important species and compared with the range of the previously described species in North America and Europe. Conodonts collected from chert and compound-type are mostly excluded in this volume, but he proposed two conodont faunas and two assemblage zones in the Japanese Permian and described 44 species and subspecies from 14 genera, among which 10 species and 1 subspecies are new to science.

This volume would be indispensable for geological and paleontological libraries as well as workers who may have interests in conodontology and later Paleozoic to early Mesozoic historical geology.

Orders must be accompanied by remittance made payable to Dr. Tsugio Shuto, Editor of the Special Papers, c/o Department of Geology, Faculty of Science, Kyushu University, Hakozaki, Fukuoka 812, Japan.

# PROCEEDINGS OF THE PALAEONTOLOGICAL SOCIETY OF JAPAN

## 学会記事

〇1982年1月22日に東京大学で開かれた評議員会に おいて賞の委員の半数改選を行ない,鎮西清高・長 谷川善和の両君がきまった。1982年度賞の委員会 は,会長のほか,高柳洋吉,木村達明(幹事)お よび上記の2名,計5名で構成される。

○同評議員会において,次の諸君の入退会,およ び特別委員への推薦が認められた(敬称略)。

[入会者] 竹村厚司(1981年度より),中川登美雄, 金 恒黙, 吉田 靖, 伊藤富夫,西田治文,松岡 篤,中川 洋,於保幸正,大矢 仁,中西ゆかり, 加藤 進,高木俊男,須田章裕,坪内弘道,国府 田良樹,間島隆一,坂本 治(18名)。

[退会者] 木村方一,石川 正,三谷勝利,平川 昌登,長瀬和雄(5名)。出光石油開発株式会社(賛 助会員)。

[新特別会員] 青島睦治, 茨木雅子, 井上英二, 大木公彦, 小笠原憲四郎, 岡村 真, 田沢純一, 野原朝秀, 松川正樹 (9名)。

○1981年中の逝去会員は, 奥津春生, 中野光雄の 両君であった。

○1982年1月23日現在の会員数は、名誉会員10,特別会員193,普通会員394,在外会員45,賛助会員9,計651名である。

○1981年9月に行なわれた地学関係5学会の協議の 結果,昭和57年度の科学研究費地質学分科の配分 委員候補として,層位古生物学1段委員に,増田孝 一郎(本会推薦),地質学一般1段委員に村田正文 (本会推薦),歌田 実の諸君が推薦された。なお, 2段委員として棚井敏雅,層位古生物1段委員とし て小畠郁生,沖村雄二,地質学一般に田口一雄の 諸君が留任している。

○1981年度日本古生物学会論文賞は、加瀬友喜君の "Early Cretaceous Gastropoda from the Choshi District, Chiba Prefecture, Central Japan (報告・記事118号)に、学術賞は、棚部一成君の"白亜紀アンモナイトの研究"、および松丸国照君の"第三紀大型有孔虫の研究"にそれぞれ 授与された。

## 日本古生物学会1982年年会 • 総会

日本古生物学会年会・総会が1982年1月23,24日 に,千葉大学教育学部を会場として開催された(参 加者222名)。

#### 海外学術集会出席報告

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# 個人講演

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の進化に関する考察竹村厚司・中世古幸次郎
Stylosphaerids の殻構造について
関東山地奥秩父大ガマタ層の微化石と地質時代
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On the Upper Permian Palaeofusulina-Cola-
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有孔虫群宇田泰子 · 斎藤常正
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Saito, T., Kadar, D., Niitsuma, N.,
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父島の始新世浮遊性有孔虫高柳洋吉・安田尚登
父島の始新世大形有孔虫松丸国照・高柳洋吉

_	父島の始新世石灰質ナンノプランクトン
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<b>.</b>	Internal structure of cassiduline Foraminifera
	Nomura R
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ついて .....谷 祖剛 (Gu Zugang)

#### 1981年度日本古生物学会論文賞推薦文

加瀨友喜君: Early Cretaceous Gastropoda from the Choshi District, Chiba Prefecture, Central Japan. Trans. Proc. Palaeont. Soc. Japan, N.S., No. 118, pp. 291-324, pls. 34-36, 1980.

本論文は千葉県銚子半島の下部白亜系銚子層群 より産した腹足類化石に関する2名の共著者による 研究成果で,加瀬君がその古生物学的な記載を行っ たものである。

白亜紀中葉は世界的に 腹足類の大きな変革期に あたり,その前後の 腹足類群の 構成には世界中の 多くの研究者が深い関心を示している。 この分類 群の進化史を理解する上に 標準地域から遠く 離れ た我国の資料は重要な役割を果すと考えられる。 我国の白亜系には多様な腹足類化石が含まれるが, 多くの場合保存状態が 十分でなく,分類上の位置 を決定するための 重要な形質(例えば 殻口部の形 状)を把握できないことが少くない。特に 海生の 腹足類については,Nerinea 類を除けば,戦前の 矢部長克・長尾 巧・松本達郎三氏の研究以後ほ とんど発展が見られなかった。

今回加瀬君によって 記載された 標本は我国の下 部白亜系としては稀に見る良好な保存状態を示し, 殻口部の 形状はもとより, 微細な彫刻や棘がよく 観察できる。 同君はこの 恵まれた 資料に基づいて 世界的な視野から,分類学的研究を進め, Nerithopsidae の1新属2新種や特徴ある Aporrhaidae の 3 新種を 含む11属12種の 腹足類化石の 分類上の位置を決定し、克明な記載を行った。特 に軸唇とふたの構造に着目した Neritopsidae の 記載分類は本科の 系統進化を考える上に重要な貢 献と考えられる。この研究は1地方化石群の記載 を主体としているが、 全く未知の 腹足類群であっ ただけに 鑑定には多大の 努力がなされたと 考えら れ, 同様の産出が 期待される日本の他地域の海成 下部白亜系の 古生物の 研究に 福音をもたらすもの である。

加瀬君の研究はまだ緒についたばかりであるが, 従来ややもすれば 等閑に 付されていた 我国の中生 代腹足類の 本格的な研究の 第一歩を開くものと評 価できる。よって 日本古生物学会は 加瀬友喜君に 今後の一層の研鑚を期待して論文賞を贈る。

#### 1981年度日本古生物学会学術賞推薦文

#### 棚部一成君:白亜紀アンモナイトの研究

棚部君は アンモナイトの パレオバイオロジーを 志し,まず北海道産異 常巻アンモナイト・スカフ イテス類の 殻特性の 個体発生的アロメトリーを解 析し,個体発生的変異と 種内変異の研究に個体群 概念を併せて 機能形態学的アプローチを示し,生 活様式の変化との関連を議論した(1975, 1977)。

同君および一部協力者との共同研究の業績は数 多くあるが,なかでも注目すべき成果を若干挙げ てみよう。アンモナイトの生活様式は世界の研究 者の間で,浮遊性あるいは底生性と長らく論議の 分かれるところであったが,さきの機能形態学的 研究ならびに北海道のチューロニアンの浅海ない し中間相における群集古生態学的解析により,連 室細管の比較強度の調査と併せて,多種類の分布・ 産状・保存状態に関する数量的資料から,白亜紀 アンモナイトの生活様式は底生性から遊泳性まで 多様であることを推論した(1978,1979)。かつ, 平滑・弱装飾型は異常巻型や装飾型のものに比し 深い環境にあることを実証し,アンモナイト群集 の生態分布モデルを提案した。

アンモナイトの発生については,現生頭足類と 異なり 孵化後幼生期と 変態をへて成長するという 説がかつて支配的であった。 棚部君は 共同研究者 とともに、 多数の 属種につき 幼期の内部構造の特 性と大分類との関連を明らかにし(1979),つづい て現生頭足類の初期の 殻形成のデータと 比較し, 微細構造や組成から, アンモナイトの 直接発達を 示すモデルを提案した(1980)。また,異常巻類の 初期の 殻形 態を 考究し,異常 巻類の 孵化 直後の "正常型" 幼体は一時的な浮遊生活を営み, 以後底 生型の生活に適応したと論じた(1981)。さらに, アンモナイトの 連室細管壁ならびに 隔壁襟の 微細 構造と組成を調べ,現生および 化石頭足類のそれ らと比較し, 隔室内液体の 移動経路についてのモ デルを提唱した(1980, 1981)。この種の研究は, 現生有殻頭足類の隔室内液体移動の 滲透圧機構に ついて 示唆を与えるばかりでなく, 頭足類進化史 を適応戦略の変化として 理解する上で,将来の貢 献が期待される。

以上を要するに,棚部君は,白亜紀アンモナイ トの進化古生物学に大きく寄与し,国の内外で注 目されるところとなった。日本古生物学会はここ に学術賞を贈り,同君の今後の研究のいっそうの 発展を期待する。

#### 松丸国照君:第三紀大型有孔虫の研究

日本における 第三紀大型有孔虫化石の 研究は, 歴史的にかなり古く、20世紀初頭の矢部長克博士 による orbitoids の記載に始まる。1920年代に至 って 矢部博士と 半沢正四郎博士の多くの 共著論文 が発表され、有孔虫研究史上、矢部・半沢時代と も称すべき活動期に入った。その後、半沢博士を 主とする大型有孔虫の研究は1960年代まで続いた が, 日本列島はもちろん広く東南アジア・太平洋 諸島などに関する成果は、 今日でも 国際的に高く 評価されている。このような歴史的背景のもとに, 群馬県富岡地域の Nephrolepidina の研究(1967) より始めた松丸君は、本州各地から産する本属の 詳細な記載・分類を行い,時空的分布と系統発生 の傾向を論ずる好著を1971年にまとめた。それ以 後, 日本列島ならびに 周辺諸島の既知の産地のみ ならず, 新産地における 大型有孔虫の産出層準を 確認し、層位的分布を明らかにする努力を続けた。

このような研究の過程で必然的に国外の特に熱 帯亜熱帯地域の有孔虫群との比較研究が必要とな るが,このことに早くより関心を抱いた松丸君は, 台湾・北米・ニュージーランド産の化石の研究 (1968—1971)を行っている。さらに橋本互君を中 心とするフィリピン・インドネシア等の地質古生 物の研究グループに1973年より参加し,主として 第三系の大型有孔虫の分類学的研究を担当して, このグループの代表作である「フィリピン産大型 有孔虫」と題するシリーズをはじめ,多くの共著 論文を発表した。

松丸君の 研究対象は 地理的に広範囲に及び、か つ研究活動ははなはだ精力的であって,1971~1980 年の間に大型有孔虫に関するものだけでも 30 篇を 超えている。 いうまでもなく, このような 成果に より西太平洋および 東南アジア地域の 大型有孔虫 に関する多くの貴重な知見がもたらされた。 しか し、もう一つ注目されるのは、 公表論文中に 海外 での知見を日本列島の化石の時空的分布の再評価 へ 還元させる 傾向のものが 見出せることである。 「日本の新生代大型有孔虫群(第1部)」(1980)は まだ 化石層位学的問題の 一部を取り上げているに すぎないが、今後豊富な資料を用いて、多様な分 類学的手法・層位学的手法等を 駆使して, より厳 密な吟味が 行われれば, 新生代大型有孔虫学への 大いなる寄与となると思われる。 よって日本古生 物学会はここに 学術賞を 贈り、 今後のいっそうの 発展を期待するものである。

## 日本古生物学会特別号 No. 24

日本古生物学会特別号 No. 24 IGO, Hisaharu: Permian Conodont Biostratigraphy in Japan (50 ページ,12図版,1981年12月15日発行)が出版されました(定価3900円,送料300円)。購入申込は特別号の 他の号と同じく日本古生物学会特別号編集委員会(福岡市東区箱崎九州大学理学部地質学教室内)(送金先: 三和銀行福岡支店普通預金口座12172;振替口座,福岡19014)にお願いします。郵送によらない直接販売は 東京大学総合研究資料館(速水格)および国立科学博物館分館(藤山家徳)でも取扱います。なお,特別号 バックナンバーの表題・販売価格(国内)・郵送料は本誌121号83ページに示してあります(残部僅少の号も ありますので売切れの節は御容赦下さい)。

## 第6回 パリノロジー国際会議(予告)

4年毎に開催されるパリノロジー国際会議 (IPC) (International Palynological Conference) の第6回 大会の第1回予告が到着しましたのでお知らせします。

開催場所: Calgary University, Calgary, Canada

期間:1984 • 8 • 24-30

IPC の講演区分・提出論文締切日等については1982年11月発行の第2回予告で示される。講演公式用語は英・仏語。学会に際しては内容要約・野外見学旅行案内書等が刊行される予定。野外見学旅行については下記があげられています。

1. Calgary 中心……Rocky Mountain, Banff 等

2. 航空機利用 ·······Hay River, Norman Well 等

3. 計画中……Nova Scotia, Quebec 等

なお出席希望者の学会内容についてのアンケートなどが付いております。この予告入手希望の向は下記に御 連絡下さい。

103 東京都中央区日本橋室町 3-2 C Mビル パリノ・サーヴエイ KK 徳永重元

(Tel 03-241-4566)

第3回太平洋地域第三系国際会議

第15回太平洋学術会議のセクションB固体地球科学と共催で開かれる。

開催場所: Dunedin University, Dunedin, New Zealand

期間:1983年2月2-4日,7日

すべての参加希望者は Pacific Science Congress, P.O. Box 6063, Dunedin, New Zealand に登録されたい。

テーマ:1) 太平洋地域の年代層序体系と対比,2) 地磁気層序と放射年代,3) 生層序,4) 古海洋と 古環境,5) 新第三紀の地球的事件(海水準,同位体他),6) 新第三系の諸境界

巡検:会期中2日間 South Island の第三系(130米ドル),会期後6日間 第三系(700米ドル) 国内問合せ先:東北大学理学部地質古生物学教室 高柳洋吉

## 環太平洋ジュラ系第1回 Fiele Conference

環太平洋地域のジュラ系に関する共同研究が IGCP の正式プロジェクトとして認められ、 第1回 Field Conference が行われる。

開催場所: Calgary University, Canada

期間: 1982年8月9日-14日

なおこの現地討論会はモントリオールで計画されている 3rd North American Paleontological Convention の後に接続されるよう計画されている。次の11のトピックに分れて討論される。

A) Ocean floor, continental margins and exotic terrains, B) Palaeogeography of sedimentary basins, C) Eustatic events and their correlations, D) Ocean currents and climatic events, E) Floral biostratigraphy, G) Macrofaunal biostratigraphy, H) Supra-regional standard zones, I) Biogeography

連絡先: Prof. G.E.G. Westermann, Dept of Geology, McMaster University, Hamilton, Ontario, Canada.

国内連絡先: 305 茨城県新治郡桜村 筑波大学地球科学系 佐藤 正

行事予定

	開催地	開催日	講演申込締切
第129回 例 会	北海道大学	1982年6月26日	1982年5月15日
第130回 例 会	三重大学	1982年10月17日	1982年8月17日
1983年年会·総会	東京大学	1983年1月22•23日	1982年11月22日

|講演申込先:〒113 東京都文京区弥生 2-4-16 日本学会事務センター 日本古生物学会行事係

### お知らせ

○本会名誉会員藤本治義君は昭和57年3月31日に逝去されました。ここに本会の発展に尽された功績を 偲び, 謹んで哀悼の意を表します。

日本古生物学会

- ○日本古生物学会では過去5年間経費の節減をはかりながら皆様の会費を据置いてきましたが、昨今の 諸物価の高騰により次第に財政が苦しくなって参りました。現在の規模の活動を続けていくには来年 度は会費の値上げをお願いしなければならなくなる見通しですので、お含みいただくと共に会費の完 納に御協力下さるようお願いいたします。
- ○日本古生物学会では年会・例会をより魅力あるものにするため、シンポジウムをはじめ各種の新しい 形の会合(例えば、テクニックの会合、ポスターセッション、ワークショップ、夜間小集会など)の 提案を歓迎します。よい企画がありましたら早目に常務委員会あてお申出下さるようお願いします。

○日本古生物学会刊行の和文誌「化石」を次のように充実させる方向で検討しています。1) B5版の 定期的(年2回発行)の刊行物とする[昭和57年度より実施]。2)現在の予約購読制を改め,誌代を 会費に含め全会員に配布する。3)原著論文を掲載し,内容を一層充実させる。これらの実施につき ましては,印刷実費に相当する会費の値上げが伴ないますので,御意見などがありましたら,化石編 集部または常務委員会まで早日にお申出下さい。

○1981年度に投稿原稿の校閲者として尽力された諸兄に感謝いたします。

○本年度より賞の委員会委員の半数改選にともない幹事が交代しました。1982年度中の各種の賞に関する問合せ,推薦依頼は鎮西清高(東京大学理学部地質学教室)にお願いします。その他の委員の役割 分担と連絡先(本誌122号参照)には変更ありません。

常務委員会

◎ 文部省科学研究費補助金(研究成果刊行費)による。

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